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# Thunderstorms and upper troposphere chemistry during the early stages of the 2006 North American Monsoon

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## Abstract

In this study, the Weather Research and Forecasting model coupled with Chemistry (WRF-Chem) is applied at 4 km horizontal grid spacing to study the meteorology and chemistry over the continental US and Northern Mexico region for the 15 July to 7 August 2006 period, which coincides with the early stages of the North American Monsoon. Evaluation of model results shows that WRF-Chem reasonably represents the large-scale meteorology and strong convective storms, but tends to overestimate weak convection. In the upper troposphere, the WRF-Chem model predicts ozone and carbon monoxide (CO) to within 10–20 % of aircraft and sonde measurements. However, the frequency distribution from satellite data indicates that WRF-Chem is lofting too much CO from the boundary layer (BL). Because ozone mixing ratios agree well with these same satellite data, it suggests that chemical production of O<sub>3</sub> in the model is overpredicted and compensates for the excess convective lofting of BL air. Analysis of different geographic regions (West Coast, Rocky Mountains, Central Plains, Midwest, and Gulf Coast) reveals that much of the convective transport occurs in the Rocky Mountains, while much of the UT ozone chemical production occurs over the Gulf Coast and Midwest regions where both CO and volatile organic compounds (VOCs) are abundant in the upper troposphere and promote the production of peroxy radicals. In all regions most of the ozone chemical production occurs within 24 h of the air being lofted from the boundary layer. In addition, analysis of the anticyclone and adjacent air indicates that ozone mixing ratios within the anticyclone region associated with the North American Monsoon and just outside the anticyclone are similar. Increases of O<sub>3</sub> within the anticyclone are strongly coincident with entrainment of stratospheric air into the anticyclone, but also are from in situ O<sub>3</sub> chemical production. In situ O<sub>3</sub> production is up to 17 % greater within the anticyclone than just outside the anticyclone when the anticyclone is over the Southern US indicating that the enhancement of O<sub>3</sub> is most pronounced over regions with abundant VOCs.

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## 1 Introduction

Convection transports chemical constituents from the boundary layer to the upper troposphere (Chatfield and Crutzen, 1984; Dickerson et al., 1987), where these chemical species undergo long-range transport and have increased lifetimes due to the cold temperatures. The global-scale impact of deep convection on the upper troposphere (UT) composition has been estimated using 3-dimensional, global-scale chemistry transport models (Lelieveld and Crutzen, 1994; Lawrence et al., 2003), which show that upper tropospheric ozone ( $O_3$ ) can be locally enhanced by upward transport in convection over polluted regions. Lawrence et al. (2003) found that the net global effect of convective transport of  $O_3$  and its precursors is to increase the amount of UT  $O_3$  by 12%. While these estimates on the importance of deep convection and  $O_3$  provide an important context for convection and chemistry studies, they are based on results from models with coarse resolution and parameterized convection. As these parameterizations are highly idealized representations of convective processes, there is concern as to the possibility of substantial uncertainties in the global estimates of gas and aerosol transport by convection.

Park et al. (2004) investigated the role of convection on  $O_3$  and carbon monoxide (CO) and on the impact of their intercontinental transport. The importance of adequately resolving mesoscale flow features for the purpose of quantifying convective transport was illustrated by contrasting simulation results produced using a uniform computation grid of  $2.0^\circ$  longitude  $\times$   $2.5^\circ$  latitude to results produced using a “stretched grid” with a  $0.5^\circ$  longitude  $\times$   $0.5^\circ$  latitude mesh over the Central US where convection is occurring. Their results showed  $O_3$  production rates to be nearly 3 times greater over Eastern North America at a 10.6 km altitude two days after convection occurred in the Central US in the lower resolution simulations ( $2.0^\circ \times 2.5^\circ$ ) than in the stretched grid simulations ( $0.5^\circ \times 0.5^\circ$ ) because of the more rapid dilution of the convective outflow plume. Thus, adequate representation of the mesoscale features is required to accurately account for convective transport. While the Park et al. (2004) study indicates that

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thunderstorms and chemistry are responsible for increased ozone in the UT, they do not point to specific geographic regions that may be affecting the UT ozone chemistry more than other regions. In this paper, we analyze results from the Weather Research and Forecasting (WRF) model coupled with Chemistry (WRF-Chem) to determine which geographic regions contribute to convective transport and to O<sub>3</sub> production in UT convective outflow regions.

Recent papers (Zhang et al., 2003; Li et al., 2005; Cooper et al., 2006, 2007, 2009) have documented the existence of an enhancement in upper troposphere ozone during July and/or August. For example, observations (interpolated from ozonesondes and Measurement of OZone and water vapour by Airbus In-service Aircraft (MOZAIC) data during August 2006) show an UT O<sub>3</sub> enhancement over the Southeast US (Cooper et al., 2007). The enhanced O<sub>3</sub> in the UT is a recurring event (Cooper et al., 2007) and is related to the anticyclone that is set up during the North American Monsoon (NAM), a seasonal shift in winds that allows low level moisture to flow into the Southwest US from the Gulfs of California and Mexico. The enhanced O<sub>3</sub> in the UT is also found in satellite data (A. M. Eldering, personal communication, 2011) and in several modeling studies (Zhang et al., 2003; Li et al., 2005; Choi et al., 2009).

It is hypothesized that the enhanced UT ozone is a result of thunderstorms and chemistry occurring during the North American Monsoon when the UT anticyclone traps the air (Li et al., 2005; Cooper et al., 2007). Thunderstorms transport boundary layer air that is rich in volatile organic compounds (VOCs) and are a source for lightning-produced nitrogen oxides (NO<sub>x</sub>). Both VOCs and NO<sub>x</sub> along with solar radiation are ingredients for producing O<sub>3</sub>. Lagrangian model simulations of a lightning-NO<sub>x</sub> tracer show that the locations of chemically produced UT O<sub>3</sub> would primarily occur in the Southern US (Cooper et al., 2009). In addition to a local chemical source of O<sub>3</sub>, high O<sub>3</sub> mixing ratios could be a result of transport of O<sub>3</sub> from the stratosphere to the UT. Simulations using the GEOS-Chem model indicate that, when the stratospheric O<sub>3</sub> is well simulated, the agreement between model results and ozonesondes is within 10 %

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in the UT (Jourdain et al., 2010), indicating that stratospheric O<sub>3</sub> may be contributing to UT O<sub>3</sub> mixing ratios.

In these previous model studies, the ability to represent the convection was limited because the models' horizontal grid spacing were too coarse to resolve convective transport. In these studies, convective parameterizations were used to represent the convective transport. Our WRF-Chem simulations are configured at 4 km grid spacing so that convective systems and convective transport are explicitly simulated. After evaluating the results of our simulation using both meteorological and chemical observation data, the WRF-Chem model results are analyzed to examine regions of the US that contribute to the UT CO and O<sub>3</sub> mass. In addition to identifying geographic regions of importance for convective transport and subsequent O<sub>3</sub> production in the UT convective outflow, the role of the anticyclone associated with the North American Monsoon on these constituents and their chemistry is examined. In particular, the contribution of in situ O<sub>3</sub> production and stratospheric air intrusions to elevated O<sub>3</sub> mixing ratios within the anticyclone are discussed.

## 2 Model description

The Weather Research and Forecasting model (Skamarock et al., 2008) coupled with Chemistry (Grell et al., 2005; Fast et al., 2006) version 3.0.1 is used to simulate the meteorology and chemistry over the US and Mexico (Fig. 1) from 00:00 UTC 10 July to 00:00 UTC 7 August 2006. There are 1200 × 900 grid points in the west-east and south-north directions, respectively, for the 4 km horizontal grid spacing domain. The vertical coordinate has 51 levels from the surface to 10 hPa. The vertical grid spacing near the surface is ~60 m, for 1–3 km altitudes it is 200–400 m, and for 5–13 km it is 540–600 m.

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## 2.1 Meteorology

Two coarse-scale WRF outer domains (36 km and 12 km grid spacing, respectively) are used to downscale the NCEP Global Forecast System (GFS) final analyses ( $1^\circ \times 1^\circ$  grid spacing) to the US and Mexico 4 km grid spacing domain shown in Fig. 1. The two outer domains, which are larger than the 4 km domain by  $\sim 5^\circ$  (12 km domain) and  $\sim 10^\circ$  (36 km domain), are run for meteorology only over a two month period using two-way nesting, and are re-initialized from NCEP meteorology approximately every week to avoid excessive drift away from observations. The results of the 12 km grid spacing domain are then used for one-way nesting onto the US and Mexico domain. For the innermost 4 km domain, the meteorology is nudged every 6 h to GFS analysis of temperature above the boundary layer and of horizontal winds throughout the model domain so that large-scale features maintain a position consistent with observations. Because the two outer domains are re-initialized approximately every week, the WRF-Chem simulation over the 4 km domain is also re-initialized for the meteorology at the same time. The chemistry is not re-initialized, but instead is continued from the previous run.

The prognostic meteorology parameters including winds, potential temperature, pressure, water vapor, and condensed water (i.e., cloud particles), tracer variables, and chemistry species are integrated forward in time using a Runge-Kutta integration method. The moisture variables, scalars, and chemistry species are advected using a monotonic scheme (Wang et al., 2009).

Cloud physics are computed with the Lin et al. (1983) single moment parameterization. This scheme predicts water vapor, cloud water, rain, cloud ice, snow, and graupel. No convective parameterization is used on the 4 km grid. The planetary boundary layer parameterization used in the simulation is the Mellor-Yamada-Janjic scheme (Janjic, 2002). At the surface, the NOAH land surface model (Chen and Dudhia, 2001) is employed. For heating rates, the Goddard scheme (Chou and Suarez, 1994) is used for short wave radiation and the Rapid Radiative Transfer model (Mlawer et al., 1997) is

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used for long wave radiation. In the US and Mexico domain, we allow the predicted aerosol distributions to affect the radiation calculation and the cloud physics as described by Chapman et al. (2009).

## 2.2 Chemistry

5 The initial and boundary conditions for the chemistry species come from 6-hourly output of the CAM-Chem model, which was run at a  $2.5^\circ \times 1.9^\circ$  resolution. Species taken from the CAM-Chem simulation include intermediate and long-lived gases (radicals such as OH and HO<sub>2</sub> are not included), dust, sea salt, and primary and secondary particulate organic carbon. Sulfate and nitrate aerosols are not included in the initial  
10 and boundary conditions.

The gas-phase chemistry mechanism used for the simulation is the RACM mechanism (Stockwell et al., 1997). A kinetic pre-processor (KPP) and Rosenbrock solver (Sandu et al., 2006) is used to solve the stiff set of gas phase chemical reactions. The MADE/SORGAM (Ackermann et al., 1998) modal approach is employed for the  
15 aerosols. Two modes (Aitken and accumulation) are solved for the following aerosol species: sulfate, nitrate, ammonium, sodium, chloride, sea salt, soil (dust), primary organic carbon, and secondary organic carbon. Dust emissions are included by assuming the dust flux ( $\text{kgm}^{-2}\text{s}^{-1}$ ) to be  $1 \times 10^{-14} \times U_*^{3.5}$  where  $U_*$  is the friction velocity. Dust emissions are also a function of the landuse type. The inorganic chemistry system  
20 is based on MARS (Saxena et al., 1986) and includes modifications suggested by Binkowski and Shankar (1995), which uses equilibrium thermodynamics to calculate the chemical composition of a sulphate/nitrate/ammonium/water aerosol.

Several emission sources are included in the simulations. The anthropogenic emissions come from the US EPA National Emission Inventory (NEI) representative of 2005  
25 typical summertime weekday (S. McKeen, personal communication, 2010) and from the Mexico NEI representative of 1999 emissions as used by Fast et al. (2009) and Hodzic et al. (2010). These emissions are reduced for Saturdays and Sundays following the findings of Stephens et al. (2008). The Saturday emissions are 75 %, 90 %, 16413

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90 %, and 93 % of weekday emissions for NO, CO, VOCs, and BC, respectively. The Sunday emissions are 60 %, 75 %, 75 %, and 80 % of weekday emissions, for NO, CO, VOCs, and BC, respectively. Biogenic emissions are calculated online using the Model of Emissions of Gases and Aerosols from Nature (MEGAN) model (Guenther et al., 2006). MEGAN uses WRF-predicted temperature and downward radiation to estimate the biogenic VOC emissions. Wildfire emissions are based on fire locations as seen by the MODIS sensors aboard the NASA Terra and Aqua satellites (Wiedinmyer et al., 2006). The injection height of the wildfire emissions are computed using a plume-rise module (Freitas et al., 2005, 2009). Aircraft emissions of CO, NO, and SO<sub>2</sub> from scheduled, charter, general aviation and military traffic for 1999 are included at a 1° × 1° resolution (Baughcum et al., 1996; 1998; Mortlock and Alstyne, 1998; Sutkus et al., 2001).

The photolysis rates are calculated using fast-TUV (Tie et al., 2003) which adjusts the rate according to the presence of aerosols and/or clouds in the grid cell. In determining the photolysis rates, the overhead O<sub>3</sub> column is set to 300 Dobson units over the US and 280 Dobson units over Mexico. Dry deposition of gases and aerosols follow the Wesely (1989) resistance method. Wet deposition of soluble gases and aerosols is calculated as described by Easter et al. (2004).

The production of NO from lightning has been implemented in this simulation. The number of lightning flashes is computed based on the maximum updraft speed (Price and Rind, 1992) within each parallel-processing tile (200 × 116 km<sup>2</sup> region) of the domain. The intra-cloud to cloud-to-ground flash ratio is based on climatology (Boccippio et al., 2001), which shows high ratios (>3) for the region between 105° W and 90° W and lower ratios elsewhere. The NO source from lightning is placed within the 20 dBZ volume of the storm, following the vertical distribution curves given by DeCaria et al. (2005). For both intra-cloud and cloud-to-ground flashes 330 moles of NO are produced per flash. Included in the simulation are two passive lightning NO<sub>x</sub> tracers, one tracing NO<sub>x</sub> produced from cloud-to-ground flashes and the second tracing NO<sub>x</sub> from intracloud flashes. Both tracers have the same source strength determined from

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the lightning-NO<sub>x</sub> parameterization, but then are only transported in the model domain; there are no sinks for these two tracers.

## 2.3 Tracers and other diagnostic parameters

To track boundary layer air, stratospheric air, and air from the horizontal boundaries, six tracers are included. The boundary layer tracers are defined as having a value of one from the surface to the planetary boundary layer (PBL) height. These tracers are maintained at the value of one in this region throughout the simulation. The stratosphere tracers are defined as having a value of one from the height of the tropopause to the top of the domain (10 hPa). Initially, for the 10–24 July segment of the simulation, the tropopause is defined as the height of the minimum temperature, which may misrepresent the actual tropopause especially when a double tropopause is present. To use a better diagnostic, the WMO definition of the tropopause is used for the 25 July–7 August segment of the simulation. These tracers also retain this value of one within the stratosphere. The horizontal boundary tracers are initially set to zero within the model domain. Their value at the horizontal boundaries is set to one throughout the simulation. Thus, the horizontal boundary tracer allows us to track air that has originated outside the model domain.

Two types of tracers are used. A non-decaying tracer ( $c_0$ ) is only transported, while its twin tracer ( $c$ ) decays with a time scale of one day. The decaying tracer loss term is

$$\frac{dc}{dt} = -kc \quad (1)$$

where  $k$  is the decay rate of  $1 \text{ day}^{-1}$ . By taking the logarithm of the ratio of the decaying tracer to the passive tracer, the age of air for the tracer can be estimated. Specifically, upon integrating Eq. (1), the age is defined as

$$t_{\text{age}} = -\frac{1}{k} \ln \left( \frac{c}{c_0} \right) \quad (2)$$

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Estimating the age of air using this method gives an idea of how long it has been since the air parcel has been lofted from the boundary layer. However, it is quite possible erroneous ages are calculated when mixing of fresh and older air occurs.

Two wildfire tracers, one passive and one with a one-day first-order decay rate, are included to follow air influenced by wildfire emissions. The wildfire tracers have an emission source strength that is the same as the CO emissions from the wildfires. As with the other tracers, there are no physical sinks (e.g., dry or wet deposition) for the wildfire tracers.

Diagnostics for CO and O<sub>3</sub> tendencies are included in the model calculation. These tendencies include the change in the CO or O<sub>3</sub> mixing ratio from advection (in all 3 directions), from the vertical mixing, from horizontal mixing, and from chemistry. These diagnostics are assigned the tendency term in the case of advection and mixing, and are calculated as the difference of the species mixing ratio before and after the chemistry solver.

### 3 Results

We have performed a WRF-Chem model simulation from 10 July to 7 August 2006, for which the 15 July to 7 August period has been analyzed to avoid the influence of the model spin up during the first few days. Thunderstorms were occurring daily over the Western Rockies and along the Gulf Coast during this period. We define four meteorological periods during the simulation to examine the composition and chemistry of the upper troposphere in relation to thunderstorm activity. Figure 2 (left panels) shows the ECMWF Interim Reanalysis (ERA-Interim) of the 300 hPa geopotential heights for one day from each of the four meteorological periods. From 15–17 July 2006 (period 1) the Northern Great Plains of the US began to experience heat wave conditions. This region was located below an anticyclone (marked by the geopotential height of ~9800 m) that was present at both 500 and 300 hPa. The anticyclone migrated westward and was over the Western US from 18–25 July (period 2) when California experienced

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anomalously high temperatures. Due to a cold front that traveled from the Great Lakes region to the Gulf of Mexico on 21–22 July, the anticyclone remained over the Western US for a full week. From 26 to 31 July (period 3) the anticyclone moved eastward, once again creating heat wave conditions over the Central Plains of the US. The anticyclone continued to move eastward and was located over the southeast and south-central region of the US from 1–7 August 2006 (period 4).

### 3.1 Evaluation of the predicted meteorology

To evaluate the WRF-Chem results, four specific dates, 18 July, 23 July, 27 July, and 1 August, at 00:00 UTC are used to illustrate how well the modeled variables compare with observations or ERA-Interim analysis. In the comparison of the 300 hPa heights from the WRF simulation with the ERA-Interim analysis, it is apparent that the simulation predicts the location and magnitude of the geopotential heights (Fig. 2) fairly well. In particular, the major features are reasonably captured. The ERA 2-m temperature maps (not shown) have regions of high temperatures ( $>310$  K or  $37^{\circ}\text{C}$ ) coincident with the high 300 hPa heights. During the early period of the simulation, temperatures are high over the Central Plains. These high temperatures are seen primarily over the west on the 23 July and then move eastward late in July. The WRF simulation largely reproduces these patterns although some regional differences in temperatures exist.

Deep convection occurred throughout the time period of the simulation. The WSI (Weather Services International) radar composites (i.e. maximum reflectivity in a column) from the National Weather Service radar network show scattered convection over the Rocky Mountain region (Fig. 3), and often show frontal or deep convective systems occurring in the Central Plains or Eastern US. The model-derived composite radar reflectivity captures these main features. Early in the simulation period on 18 July at 00:00 UTC (which corresponds to 5 p.m. local time on the West Coast and 8 p.m. local time on the East Coast), both the observations and model results show (Fig. 3a, b) scattered convection over the Rocky Mountain States (Utah, Colorado, Arizona, New Mexico) and more organized convection extending from Iowa to Lake Huron. There is

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also convection occurring in Alabama and Florida that WRF places in Georgia. The radar reflectivity on 23 July (Fig. 3c, d) shows a cold front extending from North Carolina along the Gulf Coast and into Western Texas. The WRF simulation also depicts this cold front but its location is further north than the observations in Texas and it extends further west. On 27 July the deep convection occurred in a mostly non-organized fashion over most of the US (Fig. 3e, f). The WRF results also show a lot of small unorganized, scattered convection, but are not in good agreement with observations. Because WRF is configured with a 4 km grid spacing without a convective parameterization, scattered convection, with narrow updrafts, is not expected to be represented well. On 1 August a cold front with deep convection extended from Lake Superior to New Mexico (Fig. 3g, h). The WRF results show a similar cold front, but with more numerous convective cells. In summary, the WRF simulation at grid spacing of 4 km generally reproduces the organized deep convective systems although the location of the individual storm cells is not always accurately predicted. In addition, WRF does not reproduce well the weakly organized convection.

While comparing maps of these results is a qualitative comparison, it gives insight as to how well WRF is representing the large-scale meteorology. We further analyze the reflectivity results quantitatively by comparing radar-reflectivity frequency distributions for four different regions, the Rocky Mountains, the Gulf Coast, the Central Plains, and the Midwest. These four regions are marked in Fig. 1. The frequency distributions are calculated at 8 different times of the day so that the diurnal variation can be discerned. The frequency distributions (Fig. 4) determined from the NWS radar reflectivity indicate a strong diurnal cycle over the Rocky Mountains and over the Gulf Coast with peak convection occurring at 00:00–03:00 UTC for the Rocky Mountains and 21:00–00:00 UTC for the Gulf Coast. Generally the frequency distributions from the observations peak near 20 dBZ for all four regions. The frequency distributions from the WRF simulations also have the strongest diurnal variation for the Rocky Mountains and Gulf Coast regions, but peak convection occurs at 21:00–00:00 UTC in the simulation for the Rocky Mountains and 18:00–21:00 UTC for the Gulf Coast, i.e., 3 h earlier than observations.

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Unlike the observations, the WRF results have a peak occurring at  $<10$  dBZ and then a second peak at 30–35 dBZ, where this second peak is of similar magnitude as the observed frequency of 30–35 dBZ. Realizing that convective transport will be less important for storms with reflectivity  $<20$  dBZ, we can limit the evaluation to radar reflectivity  $>20$  dBZ. Thus, we find that the frequency of deep convection is overpredicted in the Rocky Mountain region for reflectivity  $>30$  dBZ and underpredicted along the Gulf Coast for reflectivity  $>30$  dBZ, but has similar frequencies of distribution for the Central Plains and Midwest. These discrepancies are in part related to the choice of the Lin et al. (1983) microphysics scheme, which is known to create intense convective cells and weaker stratiform regions (Seifert and Weisman, 2005) because of the assumptions for producing hail. In terms of convective transport of chemical species from the PBL to the UT, it is likely that the convective transport is too large over the Rockies (where the boundary layer air ranges from fairly clean to more polluted near cities) and too small over the Gulf Coast region (where the boundary layer air is more polluted).

## 3.2 Predicted $O_3$ and CO in the upper troposphere

### 3.2.1 Spatial patterns

The modeled  $O_3$  and CO at 300 hPa ( $\sim 9.8$  km altitude) is averaged for each meteorological time period during the NAM simulation (Fig. 5). During the 15–17 July period when the anticyclone was situated over the Central Plains, low values of UT  $O_3$  (Fig. 5a) are found over the Western US. These low  $O_3$  values are not associated with active deep convection but are still correlated with the boundary layer tracer. While the Western US is susceptible to inflow from the boundary conditions, examination of model results before 15 July show convection and convective transport of high CO and low  $O_3$  occurring over Northwest Mexico that subsequently was transported west then northward to California. Over the Northern Plains, low  $O_3$  mixing ratios are also found in the UT that are associated with active deep convection in the region based on the radar reflectivity and BL tracer high mixing ratios. High values of  $O_3$  over the Central

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Plains and Midwest are correlated with the stratosphere tracer. UT CO mixing ratios are 70–90 ppbv over most of the contiguous US (Fig. 5b) during this same period. High CO mixing ratios seen over the Gulf of Mexico and Southeast US arise from air circulating around the anticyclone with an origin from a mesoscale convective system over Kansas at the beginning of the simulation (10 July). This phenomenon is seen in both the WRF-Chem simulation and in the CAM-Chem simulation that initialized and provided boundary conditions for WRF-Chem.

During the 18–25 July period, when the anticyclone was situated over the Western US, modeled UT O<sub>3</sub> mixing ratios over much of the contiguous US are 70–100 ppbv (Fig. 5c), where the lower mixing ratios are attributed to active convection transporting low O<sub>3</sub> from the boundary layer to the UT over the Northern Plains and to long range transport from the Pacific. High mixing ratios of modeled UT CO (Fig. 5d) are found along the Gulf Coast, Northern Mexico, and the Southwest US, which are regions influenced by convective transport and horizontal advection in the UT downstream of the deep convection.

During the 26–31 July period, when the anticyclone moved back over the Central US, modeled UT O<sub>3</sub> is high over the Western US and northern half of the contiguous US (Fig. 5e), where air is being rapidly transported from west to east. Thus, UT O<sub>3</sub> in the northern half of the domain represents long range transport that may be associated with stratospheric O<sub>3</sub> upstream of the model domain. In the southern half of the US, the winds at 300 hPa are much slower and circulate around the anticyclone allowing the accumulation and chemistry of O<sub>3</sub> to occur. Modeled UT CO mixing ratios are high over the Southwest US near the center of the anticyclone and over the mid-Atlantic coast (Fig. 5f). Examination of the 3-hourly model output shows that the high mixing ratios over the southwest are a result of CO accumulating in this region around the anticyclone, while the high CO off the east coast is a result of convective transport either over the Eastern US or over the Northern US that was subsequently transported eastward.

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During the 1–6 August period, the anticyclone moved to the Southeast and South-Central US. The simulated  $O_3$  at 300 hPa is found to be  $>100$  ppbv in the center of this anticyclone (Fig. 5g) as a result of air entraining into the anticyclone region and accumulating in this area. The high  $O_3$  over Canada and the Northwest US originates from the stratosphere. The modeled CO mixing ratios at 300 hPa are high over the Eastern US and low over the Western US (Fig. 5h), which is a result of clean air advecting in behind an eastward moving front.

### 3.2.2 Comparison with satellite data

Curtains of modeled  $O_3$  and CO mixing ratios are compared with observations from the Tropospheric Emission Spectrometer (TES) instrument onboard the AURA satellite. The TES is an infrared Fourier Transform Spectrometer (FTS), which measures spectral infrared (IR) radiances between  $650\text{ cm}^{-1}$  and  $3050\text{ cm}^{-1}$  in a limb viewing and a nadir (downward-looking) mode. The technique has a horizontal footprint of 5.3 km by 8.4 km in the nadir view, which provides vertical information of the measured species (e.g. Worden et al., 2004 and references therein). Under cloud-free conditions, TES is sensitive to both lower and upper tropospheric ozone with a vertical resolution between 6–8 km. TES is sensitive to CO in the middle and upper troposphere under nominal clear-sky conditions (Worden et al., 2004). The data used here were taken in the “step-and-stare” mode, in which observations are taken every 25 km over an arc of 60 degrees. To compare WRF-Chem results with the TES measurements, the model results are modified following the same procedure used in the satellite retrieval algorithm. The procedure applies the TES a priori data and averaging kernel, which designates the vertical sensitivity of the retrieval, and then removes results for regions where cloud interference was found.

As an example, the TES data from the orbit pass over the Central US on 22 July 2006 at  $\sim 19:40$  UTC are compared with WRF-Chem output at 21:00 UTC on the same day. This overpass is closest in time ( $\sim 4$  h) with the west to east frontal convection occurring in Texas (Fig. 3c). The WRF-Chem simulation captures the general features of the  $O_3$

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and CO distributions (Fig. 6). High ozone mixing ratios are seen above 600 hPa level for the 22–35° N latitude range for both the TES data and the WRF-Chem results. North of 35° N (behind the cold front) ozone mixing ratios are <50 ppbv. While the WRF-Chem results do not have ozone mixing ratios as high as those shown by TES in the UT near 30° N, the structure of the high ozone feature is quite similar to that seen in the TES data. The TES CO product shows CO mixing ratios of >100 ppbv reaching up to 300 hPa near 30° N near the frontal convection (Fig. 6d). Likewise, the WRF-Chem results show evidence of convective lofting to the upper troposphere near the frontal convection but also further south over the Gulf of Mexico (~27° N) where air from the mid-Atlantic states was lofted a day or so earlier via convection.

To quantify the WRF-Chem upper troposphere O<sub>3</sub> and CO mixing ratios with the TES data, frequency distributions are computed for the “step and stare” TES tracks passing over the model domain between 15 July and 6 August for the 200–400 hPa altitude range. For points within 100 km of the TES overpass, WRF-Chem frequency distributions were computed based on raw output and also with output adjusted to take into account the influences of the TES averaging kernel. The frequency distribution for ozone (Fig. 7a) is broad with a peak of about 75 ppbv for both TES data and WRF-Chem results. The overlapping of the WRF-Chem results and TES data indicates that WRF-Chem is representing ozone in the UT well. The CO frequency distribution (Fig. 7b) has a narrower peak than the ozone distribution. Both TES and WRF-Chem CO peaks are at 80 ppbv. However, the TES frequency distribution has CO mixing ratios ~25 % lower than that in WRF-Chem indicating that the model is lofting too much boundary layer air. Ozone shows <20 % difference in low mixing ratios suggesting that chemical production of O<sub>3</sub> in the model compensates for the excess convective lofting of BL air.

### 3.2.3 Comparison with ozonesonde data

Ozonesondes are launched from Trinidad Head, CA, Boulder, CO, and Huntsville, AL representing three different regions of the conterminous US. During July 2006, these

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balloons were launched approximately weekly and in August 2006 the balloons were launched daily as part of the special observations for the Texas Air Quality Study (Cooper et al., 2007; Thompson et al., 2008). Between 15 July and 6 August, there were 8 balloons launched at each of the locations. The data from each of these sites were averaged and standard deviations calculated. The WRF-Chem profiles were obtained, first, by averaging model result profiles over a  $48 \times 48 \text{ km}^2$  area (a region that is small enough to have somewhat homogeneous ozone mixing ratios, yet large enough to account for deviations caused by poorly-predicted small-scale meteorology) centered on the grid point that is nearest the ozone launch site. These average profiles were obtained for each of the observed times. Then, like the observations, these profiles from different days for each location were time-averaged and standard deviations calculated.

The  $\text{O}_3$  profiles (Fig. 8) at Trinidad Head show the most variance, especially in the 8–11 km altitude range. This high variability is seen only in the observations and much of the variance can be attributed to stratospheric air, as indicated by the low water vapor mixing ratios in the ozonesondes, that was not replicated by WRF-Chem because often the stratosphere-troposphere exchange occurred upstream, outside the model domain. For both Trinidad Head and Huntsville ozonesondes, the agreement between observations and model results is quite good from the surface to 10 km or higher. The average ozone profile for Boulder shows that WRF-Chem overestimates ozone in the mid- to upper troposphere compared to observations, but predicted ozone is still within 1 standard deviation of observations from the surface to 13 km altitude. Above 12 km, the WRF-Chem profiles have much higher ozone mixing ratios than observed suggesting the tropopause is too low in WRF-Chem. This may be a result of the WRF configuration for WRF version 3.0.1 where the water vapor mixing ratio in the stratosphere is prescribed to be too high in the long-wave radiation algorithm (RRTM) as noted by Cavallo et al. (2011).

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### 3.2.4 Comparison with MOZAIC data

During the 15 July to 3 August 2006 period, there were 46 profiles of MOZAIC data (Marenco et al., 1998) available within the WRF-Chem domain. Most of these profiles occurred for take-offs and landings in Atlanta, GA, Dallas, TX, and Philadelphia, PA. As with the ozonesonde data, the MOZAIC O<sub>3</sub> and CO profiles were averaged and standard deviations calculated for each location. The WRF-Chem profiles were constructed by finding the nearest model grid cell to each of the MOZAIC data points. The profiles were then averaged for each of the locations and standard deviations were calculated. These average profiles (Fig. 9) show a reasonable agreement between observations and model results between the surface and 12 km altitude for both O<sub>3</sub> and CO mixing ratios. Larger differences were found for CO below 2 km altitude with an overprediction of ~50 ppbv in modeled CO. For the 6–12 km altitude region, simulated O<sub>3</sub> is 10–20, 5–15, and 0–30 ppbv higher than observations for Atlanta, Dallas, and Philadelphia, respectively. The modeled CO is generally lower than observations in this altitude range (by <20 ppbv for Atlanta and Dallas) except for Philadelphia near 10 km altitude where WRF-Chem overpredicts CO by ~35 ppbv. All of the modeled profiles are within one standard deviation of the observations.

In summary, WRF-Chem predicts CO and O<sub>3</sub> mixing ratios in the mid- to upper troposphere to within 10–20 % of the satellite, ozonesonde, and MOZAIC measurements. On the other hand, the frequency distribution comparison between WRF-Chem and TES satellite data suggests that WRF-Chem is lofting too much CO from the boundary layer. Because ozone mixing ratios agree well with these same satellite data, it suggests that chemical production of O<sub>3</sub> in the model is overpredicted or there is too much stratosphere to troposphere transport in the model. The contribution of each of these ozone sources is discussed in the next section.

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## 4 Effect of convection on UT CO and O<sub>3</sub> – geographic region perspective

To examine the effect of convection from different geographic regions on UT CO and O<sub>3</sub>, time series of the area-averaged mass of the species and tracers and the diagnostic ozone chemical tendency term are analyzed as well as the ozone chemical production based on the age of the BL tracer in the UT region.

### 4.1 CO, O<sub>3</sub>, and tracer masses

Figure 10 shows the time evolution of the mass per km<sup>2</sup> of CO and O<sub>3</sub> in the 8–12 km altitude range during the 3-week simulation analysis period as predicted by WRF-Chem. The decaying boundary layer and stratosphere tracers are also plotted to indicate the influx of air from those two regions. The decaying boundary layer and stratosphere tracers were unfortunately re-initialized to zero when meteorology re-starts occurred on 24 July and 2 August. The decaying boundary layer tracer time series shows the daily influx of boundary layer air through vertical transport, and its decay due to the 1-day lifetime. The daily variation is most pronounced in the Rocky Mountain area. However, the convective lofting of boundary layer air is likely an overestimate based on the evaluation of radar reflectivity in the Rocky Mountain region (Sect. 3.1). The CO area-averaged mass does not show the same distinct daily variation that the decaying boundary layer tracer has, likely because CO has a chemical lifetime much longer than 1 day in the upper troposphere. During the 23 day period, the CO area-averaged mass in the upper troposphere increases slightly (10 % or less) in all regions except the Rocky Mountains, which experiences a 20 % increase for the first 17 days and then a strong decrease after the passage of the cold front on 1 August. These same trends for each region can be seen in the plots of Fig. 5.

The decaying stratosphere tracer has episodic increases associated with either stratosphere-troposphere dynamics north of the jet stream (31 July to 1 August in the West Coast region) or with cold front passages (22–24 July in the Midwest). There is a correlated increase in O<sub>3</sub> area-averaged mass for these same time periods indicating

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that the high ozone mixing ratios are from the stratosphere. During the 18–25 July period when the anticyclone was situated over the Western US, the  $O_3$  area-averaged mass in the upper troposphere increases from  $2.25 \text{ kg km}^2$  to  $2.75 \text{ kg km}^2$ , a 22 % increase, which is a result of high  $O_3$  being transported into the West Coast region from Northwest Mexico and low  $O_3$  being transported out of the West Coast region. The other regions do not show an increase in  $O_3$  area-averaged mass coincident with the anticyclone location in that region (e.g. the Central Plains  $O_3$  area-averaged mass decreases when the anticyclone is over the Central US during 15–17 July and 26–31 July). However, the  $O_3$  area-averaged mass over the Gulf Coast substantially increases ( $2.0$  to  $2.7 \text{ kg km}^2$  or 35 % increase) from 30 July to 1 August, which is a result of high ozone originally from the stratosphere moving from the Midwest (note the peak in  $O_3$  and stratosphere tracer in the Midwest on the 29 July) into the Gulf Coast region and subsequently being accumulated by the anticyclone that is in place in early August.

## 4.2 Ozone production

Figure 10 includes the time evolution of the ozone chemical tendency per  $\text{km}^2$  in the 8–12 km altitude region as predicted by WRF-Chem. The ozone chemical tendency is the difference between  $O_3$  before and after the chemical solver routine in WRF-Chem and we plot the values averaged over 3 h periods. While most of the convective transport occurs in the Rocky Mountain region (Fig. 10), most of the ozone chemical tendency occurs over the Gulf Coast and Midwest regions where VOCs other than CO and  $\text{CH}_4$  are more abundant near the surface. The West Coast shows the lowest net  $O_3$  production per area of all 5 regions.

There are apparent correlations between the tracers and the  $O_3$  production. Generally, when stratospheric air is in the region the  $O_3$  production is less than when  $O_3$  is not elevated from stratospheric influences. These instances are most prominently seen (Fig. 10) in the Midwest from 22 to 25 July and on 28 July, in the West Coast from 31 July to 2 August, and in the Rocky Mountains from 27 to 29 July. There are also



several times when the decaying boundary layer tracer and the  $O_3$  production increase together indicating more  $O_3$  production occurring in fresh boundary layer air that has been transported to the upper troposphere than in the background upper troposphere air. These instances are easily seen (Fig. 10) in the Gulf Coast region for 17–23 July and 27 July–1 August, in the Central Plains for 16–22 July and 29 July to 2 August, and in the Midwest from 20–22 July and 26–28 July. The West Coast has relatively higher  $O_3$  production in the upper troposphere soon after the anticyclone has moved from the West Coast to the Central Plains (26–27 July). The increase of  $O_3$  over the West Coast from 18–20 July is not correlated with either in situ ozone production or stratospheric air, but instead is a result of high  $O_3$  air being advected into the region from Northwestern Mexico.

### 4.3 $O_3$ production based on age of air in each region

To further understand when and where ozone is being produced in the upper troposphere, the model results for the 8–12 km altitude region are sampled based on the age of the air since it has left the boundary layer and on the geographical region. The age of air, as described in Sect. 2.3, is determined from the logarithm of the ratio of the passive to the decaying boundary layer tracer. Because mixing of new boundary layer air with old air can result in an age representing neither young nor old air, the analysis is performed using broad age categories: 0–12 h, 12–24 h, 24–48 h, 48–72 h, and 72–120 h since the air exited the boundary layer. As described above, the ozone chemical tendency and reaction rates are calculated at each grid point in the 8–12 km altitude range throughout the model domain. Within each region, the age of air is also calculated at each grid point between 8 and 12 km altitude. The rate is identified per region and age of air, summed for each category, and then divided by the number of grid points in each category to give a rate per volume result. The results are further averaged over the four meteorological time periods, 15–17 July, 18–25 July, 26–31 July, and 1–7 August, so as to examine the net ozone production in the context of the meteorological scenario.

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In a similar fashion, the rates of several oxidation reactions that form peroxy radicals (and presumably further react with NO to form NO<sub>2</sub>) are determined. In thunderstorm outflow regions, where NO generated from lightning exists, there should be sufficient NO to react with the peroxy radicals such that NO<sub>2</sub> and subsequently O<sub>3</sub> are produced.

5 While the role of lightning-NO<sub>x</sub> in these simulations will be analyzed in a future paper, evaluation of the lightning flash rate with NLDN data shows that the flash rate is under-predicted (Wong et al., in preparation). Here, we focus on the contribution of specific VOC + OH oxidation reactions to form peroxy radicals to understand where and when the oxidation occurs in the convective outflow regions. Several reactions in the RACM  
10 mechanism produce peroxy radicals. These reactions include oxidation by OH of CO, CH<sub>4</sub>, propane (HC3), ethene (ETE), terminal olefins (OLT), isoprene (ISO), formaldehyde (HCHO), hydroxy-ketone (HKET), methyl hydroperoxide (OP1), other organic peroxides (OP2), and organic nitrates (ONIT). Note that these last two are compounds representing a group of compounds and not any specific member of the group. The  
15 reaction rates were calculated using the model output of the temperature and pressure (to calculate the rate coefficient) and the mixing ratios.

The Rocky Mountain (Fig. 11) and Gulf Coast (Fig. 12) regions are used to illustrate the results, but the results are discussed for all regions. The ozone chemical tendency and the CO + OH reaction rate are normalized by their respective total values, as listed  
20 in the figure, for each region and meteorological time period. The other reaction rates of hydrocarbons with OH are normalized by the CO + OH total rate to show the relative importance of each reaction.

The analysis shows that more than half of the ozone production (e.g. Figs. 11–12) takes place in the first 24 h after convective lofting for all the geographic regions. Many  
25 of these storms occur in the late afternoon to evening (Fig. 4), thus the next day's photochemistry is responsible for the chemical production of ozone. Interestingly, the convective outflow over the Central Plains experiences most of its ozone production in air 12–24 h old; convective outflows from afternoon thunderstorms over the Rocky Mountain region move eastward into the Central Plains where oxidation occurs the

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following morning, and additionally some of these same storms propagate eastward into the Central Plains at night, but their outflows do not undergo oxidation until daylight.

The oxidation rates for all the regions and meteorological time periods show that CO + OH is the dominant pathway for producing peroxy radicals, but oxidation of organic peroxides (OP2 in Figs. 11–12), methane, HCHO, and CH<sub>3</sub>OOH (OP1 in Figs. 11–12) are important contributors to peroxy radical production. The organic peroxides are important for producing peroxy radicals, but its predicted production may not be realistic because of assumptions in the parameterized chemistry or underpredicted lightning NO<sub>x</sub>. In addition, observational studies (e.g. Snow et al., 2007) have not shown appreciable concentrations of these compounds in the upper troposphere. Thus, a careful examination of the role of organic peroxides in the upper troposphere chemistry should be pursued.

Oxidation by OH of VOCs other than CO, CH<sub>4</sub>, HCHO, and organic peroxides can be important in the Central Plains, Midwest, and Gulf Coast regions. The Central Plains and Midwest regions include oxidation of species of anthropogenic origin, such as propane. Over the Gulf Coast region, oxidation of isoprene and its products (e.g. MACR) in the convective outflow is very important (Fig. 12) for producing peroxy radicals, while isoprene oxidation has little contribution to peroxy radical formation over the Rocky Mountains.

In summary, this analysis shows that (a) more than half of the ozone production takes place in the first 24 h after convective lofting for all geographic regions, (b) organic peroxides are important for producing peroxy radicals, but its predicted production may not be realistic because of assumptions in the parameterized chemistry or possibly underpredicted lightning NO<sub>x</sub>, and (c) VOCs other than CO, CH<sub>4</sub>, HCHO, and organic peroxides are important in the Central Plains, Midwest, and Gulf Coast regions for peroxy radical production.

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## 5 Effects of convection on UT CO and O<sub>3</sub> – anticyclone perspective

To examine the processes producing and maintaining high ozone mixing ratios within the anticyclone (as observed in previous studies, e.g. Cooper et al., 2007), the model output can be sampled for air within the upper troposphere anticyclone and just outside the anticyclone. The anticyclone is defined as the region where the geopotential height is greater than 9730 m at 300 hPa, while the air just outside the anticyclone is defined as the region where the geopotential height is between 9680 and 9730 m at 300 hPa and is limited to being within 600 km of the anticyclone. For both regions, a volume of air is sampled from 8 to 12 km altitude. The mass of CO, O<sub>3</sub>, decaying boundary layer tracer, and decaying stratosphere tracer are normalized by the area of the anticyclone or the region just outside the anticyclone as defined at the 300 hPa pressure level. The anticyclone exists on all days of the simulation except on the 31 July when the anticyclone weakened and geopotential heights were always less than 9730 m.

Because the model simulation contained a passive tracer of stratospheric air instead of a stratospheric O<sub>3</sub> tracer (e.g. Roelofs and Lelieveld, 1997; Wang et al., 1998), correlations between O<sub>3</sub> and the stratosphere tracer are used to determine contributions from the stratosphere. As a result, it is not possible to quantify the contribution of various O<sub>3</sub> sources here.

### 5.1 CO, O<sub>3</sub>, and tracer masses

Figure 13 shows the mass per square kilometer of CO, O<sub>3</sub>, decaying boundary layer tracer, and decaying stratosphere tracer for the anticyclone and just outside the anticyclone. The CO and O<sub>3</sub> area-averaged mass is similar for the two regions (1.4 to 1.8 kg km<sup>2</sup> for CO and 2.0 to 3.2 kg km<sup>2</sup> for O<sub>3</sub>), although during the last week of the simulation there is somewhat more CO and O<sub>3</sub> in the anticyclone, which is located over the Southeastern and South Central US, than just outside the UT anticyclone. The decaying boundary layer and stratosphere tracers show much more variability than CO and O<sub>3</sub>. Especially in the anticyclone, there is a diurnal variability for the BL tracer

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indicating a daily lofting of the boundary layer air to the upper troposphere within the anticyclone. However, there are specific periods during which stratospheric air resides in the anticyclone (e.g. 25–30 July) indicating that while BL air is being lofted stratospheric air is also being entrained into the upper troposphere. During the 18–31 July time periods (West Coast anticyclone that then transitions to the Central US), the decaying boundary layer tracer is 30–35 % greater within the anticyclone compared to just outside the anticyclone. In contrast, the decaying stratosphere tracer is always smaller within the anticyclone compared to just outside the anticyclone.

## 5.2 Ozone production

The ozone chemical tendency per square kilometer (Fig. 13, right axis) has similar values in the two regions, peaking between 0.004 and 0.007 kg km<sup>2</sup>. On average, the ozone chemical tendency is 8 % greater within the anticyclone than just outside the anticyclone. During the 26–31 July and 1–7 August time periods, even greater ozone chemical tendencies (17 % and 10 % increase, respectively) occurred within the anticyclone. Thus, when the anticyclone transitioned from the Western to the Eastern US more ozone was chemically produced within the anticyclone, possibly explaining the observed enhancement of ozone over the Southeastern US (e.g. Cooper et al., 2007). In addition, these increases in ozone chemical tendency are in line with the results of Lawrence et al. (2003) who found a global effect of convection on O<sub>3</sub> to be a 12 % increase.

## 5.3 O<sub>3</sub> production based on age of air in each region

To further understand the ozone production within and just outside the UT anticyclone, the model results for the 8–12 km altitude region are sampled based on the age of the air since it has left the boundary layer and based on the anticyclone region. The rates are calculated in the same manner as was done for the geographical region analysis, but are normalized by the number of grid points in each category to give a rate per

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volume result. Similar to the geographical regional analysis, the ozone chemical tendency and the CO + OH reaction rate are normalized by their respective total values, as listed in the figure, for each region and meteorological time period. The other reaction rates of hydrocarbons with OH are normalized by the CO + OH total rate to show the relative importance of each reaction.

The 1–7 August time period (Fig. 14) is used to illustrate the results. For both within and just outside the anticyclone, approximately 50 % of the oxidation of CO, methane, NMHCs and other VOCs happens within the first 24 h since the air was lofted from the boundary layer. When the anticyclone is located over the Southeastern and Southern US (1–7 August), isoprene chemistry becomes important, especially in the anticyclone in which oxidation of isoprene, methacrolein, and organic peroxides are 40 %, 40 %, and 20 % greater, respectively, within the anticyclone than outside it. The oxidation of most of these hydrocarbons primarily occurs in air <24 h old. However, organic peroxide and CH<sub>3</sub>OOH oxidation have nearly equal contributions from the different ages of air, suggesting that these peroxides contribute to peroxy radical production from both lofting of these species from the boundary layer and production from other hydrocarbon oxidation reactions in a low NO<sub>x</sub> environment.

In addition to the source of O<sub>3</sub> from the chemistry, there is a source of O<sub>3</sub> from the stratosphere that is concurrent with the convection, especially during the last 14 days of the simulation (Fig. 13). Much of the stratospheric air in the upper troposphere occurs in regions marked as 3–5 days since convective lofting. For example, the instantaneous results at 21:00 UTC on 5 August (Fig. 15) show stratospheric air at 250 hPa within and along the edge of the anticyclone, as defined by the 300 hPa geopotential height of 9730 m. This same region has air that is more than 3 days old. One can also see recently lofted air often next to the older air. Plumes of O<sub>3</sub> that are correlated with the stratosphere tracer contain mixing ratios of up to 200 ppbv at 250 hPa. These plumes are in both the anticyclone and just outside the anticyclone. The high ozone mixing ratios originating from the stratosphere appear to be greater and more widespread than the O<sub>3</sub> in UT convective outflow. For example, O<sub>3</sub> is ~75 ppbv above Utah and Colorado

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(Fig. 15) where the age of the BL tracer is  $<24$  h, while  $O_3$  is  $\sim 120$  ppbv over the Kansas-Missouri border where the stratosphere tracer also exists. It is the combination of stratospheric air entraining into the anticyclone and in situ  $O_3$  chemical production that contributes to the production and maintenance of high  $O_3$  in the UT.

## 6 Conclusions

A high-resolution simulation of meteorology and chemistry was performed for the continental US and Northern Mexico region to examine the contributions of convective transport in different geographic and meteorological regions to the upper troposphere composition and chemistry. The WRF-Chem results are evaluated with ECMWF ERA-Interim geopotential height and 2-m temperature, NWS radar reflectivity composite data, TES CO and  $O_3$  satellite retrievals, ozonesonde data, and MOZAIC aircraft vertical profiles. The evaluation points out that WRF-Chem reproduces the large-scale meteorology and more intense storms ( $>35$  dBZ reflectivity) well, but tends to overestimate weak convection ( $<20$  dBZ reflectivity). The WRF-Chem model predicts upper troposphere CO to within 10–20 % of measurements and overpredicts upper troposphere  $O_3$  by 10–25 %.

An analysis of 23 simulated days is performed to understand the role of deep convection on the composition and chemistry of the upper troposphere. Analysis of different geographic regions (West Coast, Rocky Mountains, Central Plains, Midwest, and Gulf Coast) reveals that much of the convective transport occurs in the Rocky Mountains, while much of the ozone chemical production occurs in the upper troposphere over the Gulf Coast and Midwest regions where both CO and VOCs are abundant. In all regions most of the ozone chemical production occurs within 24 h since the air was lofted from the boundary layer. Besides CO, organic peroxides (excluding  $CH_3OOH$ ) produce 10–30 % of the peroxy radicals, which potentially create  $NO_2$  and  $O_3$ . Over the Gulf Coast region, in particular, biogenic VOCs (e.g. isoprene and methacrolein) are important UT ozone precursors.

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Analysis of the anticyclone associated with the North American Monsoon and of the region just outside the anticyclone shows that there are fairly similar amounts of  $O_3$  in both regions. The  $O_3$  is a product of both stratospheric air (as indicated by a stratosphere tracer) and in situ  $O_3$  chemical production (as indicated by the chemistry diagnostic). When the anticyclone is over the Southern US, there is up to 17 % more in situ  $O_3$  production within the anticyclone than just outside the anticyclone. Examination of instantaneous model results show both higher concentrations and spatially more widespread  $O_3$  co-located with the stratosphere tracer within the UT anticyclone than  $O_3$  co-located with BL air recently lofted to the UT. While both chemical and dynamical processes are enhancing the UT  $O_3$ , ozone from the stratosphere appears to have the greater contribution to this enhancement.

These results provide some insight into the chemistry occurring in convective outflow regions, but inherent in the results are uncertainties in the chemistry mechanism parameterization, emissions inventory, and the prediction and representation of small-scale convective storms. Of particular concern is the chemistry parameterization that lumps reactive organic peroxides and organic aldehydes, because previous studies have found that  $H_2O_2$  and  $CH_3OOH$  are the dominant peroxides in the upper troposphere. Thus, chemical mechanisms must continue to be improved. Field studies, such as the Deep Convective Clouds and Chemistry campaign, should provide more information on the composition and ozone production in the UT convective outflow regions, both near convection and downwind, so as to put more constraints on the UT chemistry.

**Acknowledgements.** The work was partially supported by NASA grant NNX07AM47G. The National Center for Atmospheric Research is supported by the National Science Foundation. The authors appreciate the computer time made available for this project and for the assistance of the supercomputer engineers from the CISL at NCAR and from the NASA Ames computing group. Stuart McKeen (NOAA/ESRL) is thanked for providing the 4 km gridded anthropogenic emissions from the EPA inventory. Christine Wiedinmyer and Louisa Emmons provided biogenic VOC and fire emissions data and CAM-Chem results for the boundary conditions. The Global Monitoring Division of NOAA/ESRL is kindly acknowledged for the ozonesonde data. We also acknowledge the European Commission and its partner institutions in Germany, France



and the UK for the support to the MOZAIC project. We thank Anne Boynard and Christine Wiedinmyer for helpful comments on the manuscript.

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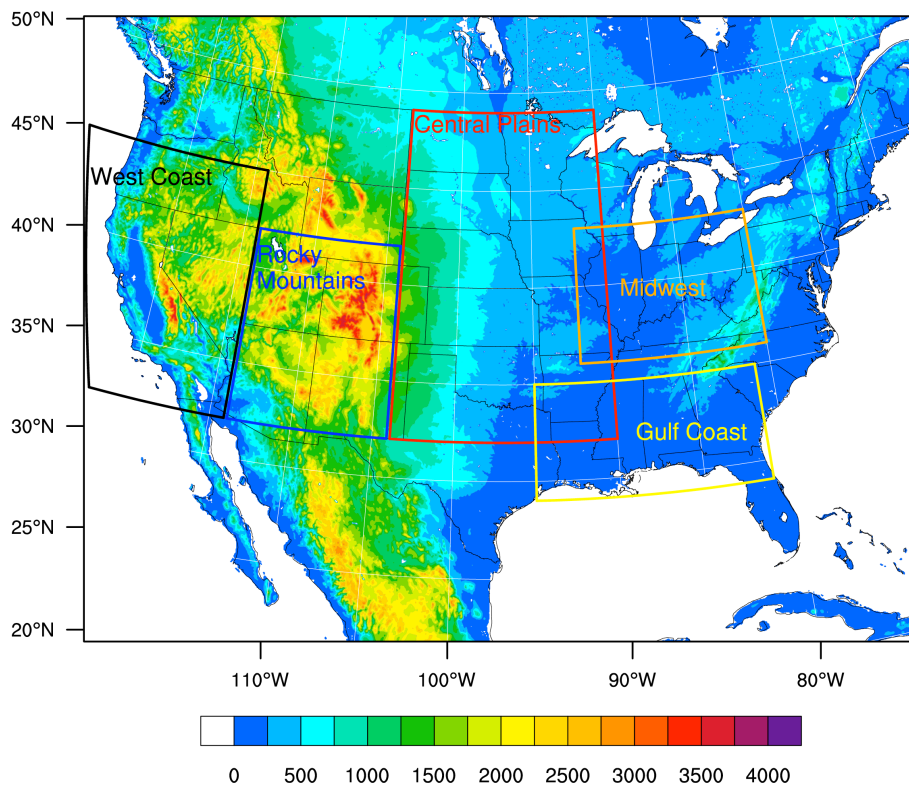
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**Fig. 1.** Domain and terrain height of the model simulation. Marked regions are used in the radar reflectivity evaluation and the chemistry analysis.

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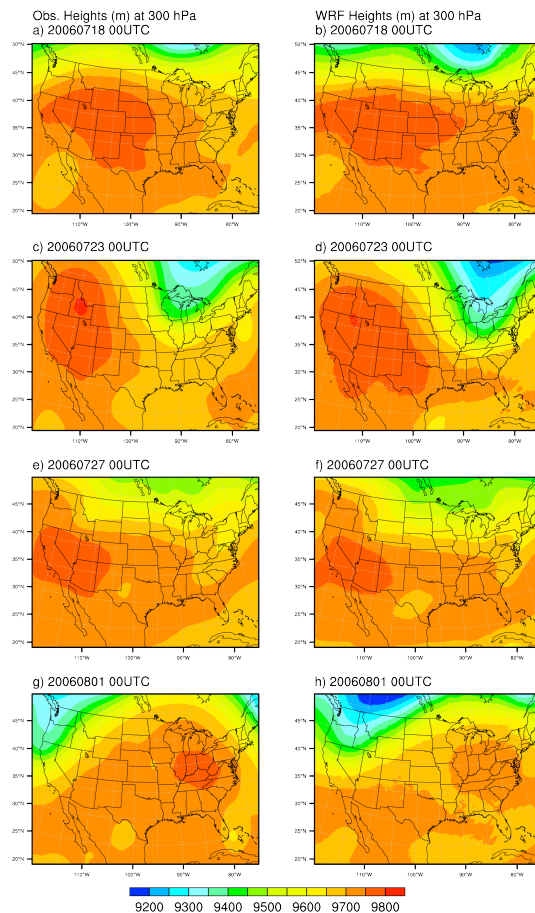
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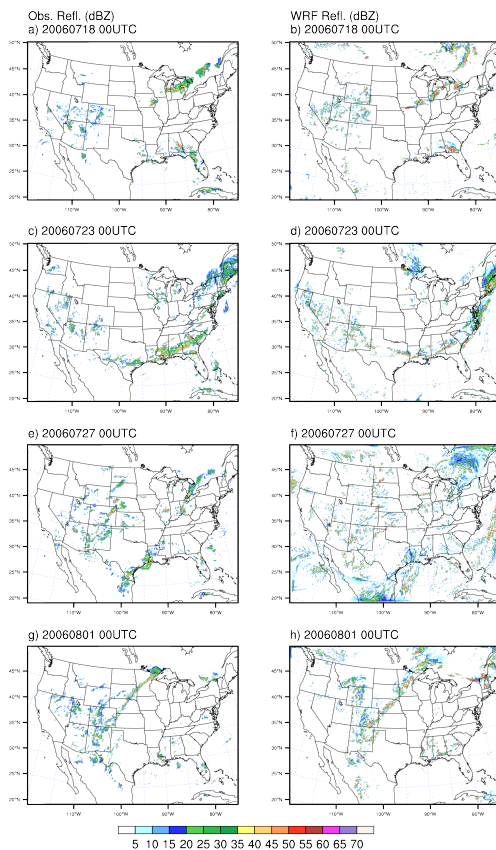
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**Fig. 2.** Spatial distribution of 300 hPa geopotential heights from the ECMWF ERA-Interim re-analysis (left panels) and WRF-Chem model results (right panels) at 00:00 UTC for 18 July 2006 (**a, b**), 23 July 2006 (**c, d**), 27 July 2006 (**e, f**), and 1 August 2006 (**g, h**).





**Fig. 3.** National Weather Service NEXRAD observed (left panels) and WRF-Chem simulated (right panels) maximum radar reflectivity at 00:00 UTC for 18 July 2006 (**a, b**), 23 July 2006 (**c, d**), 27 July 2006 (**e, f**), and 1 August 2006 (**g, h**). Note that the observed radar reflectivity is only for the US while the model-derived reflectivity includes storms in Canada and Mexico.

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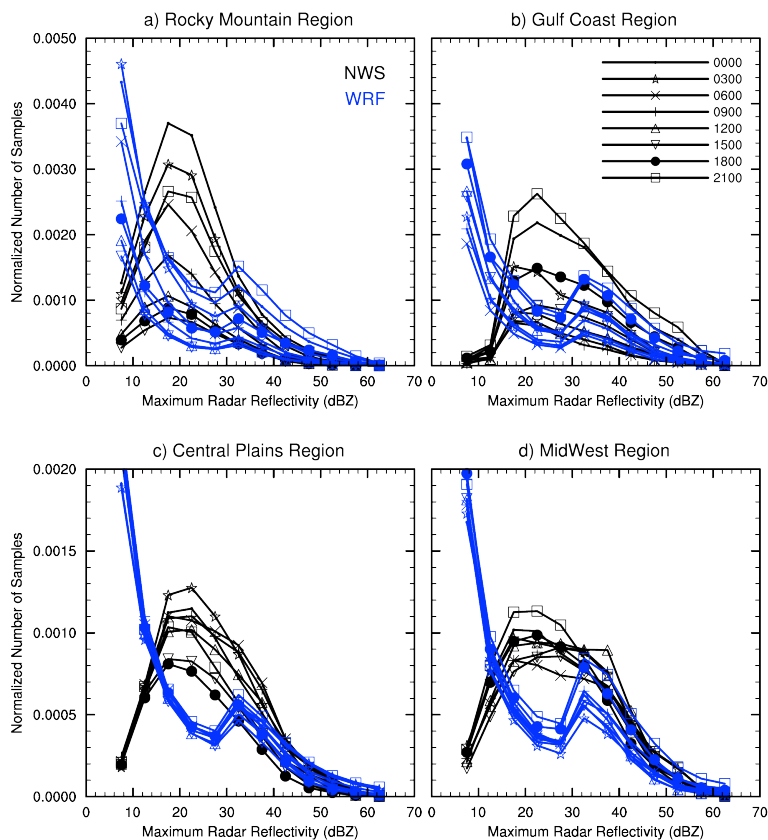
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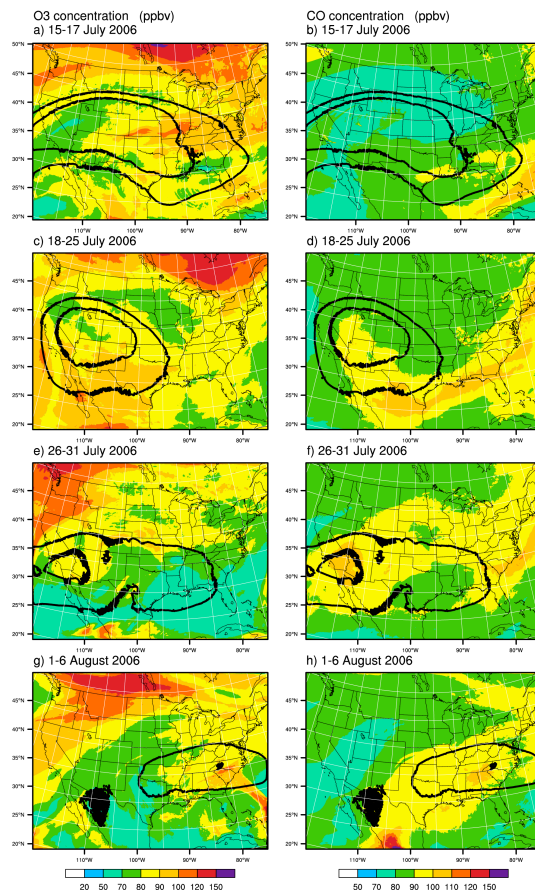
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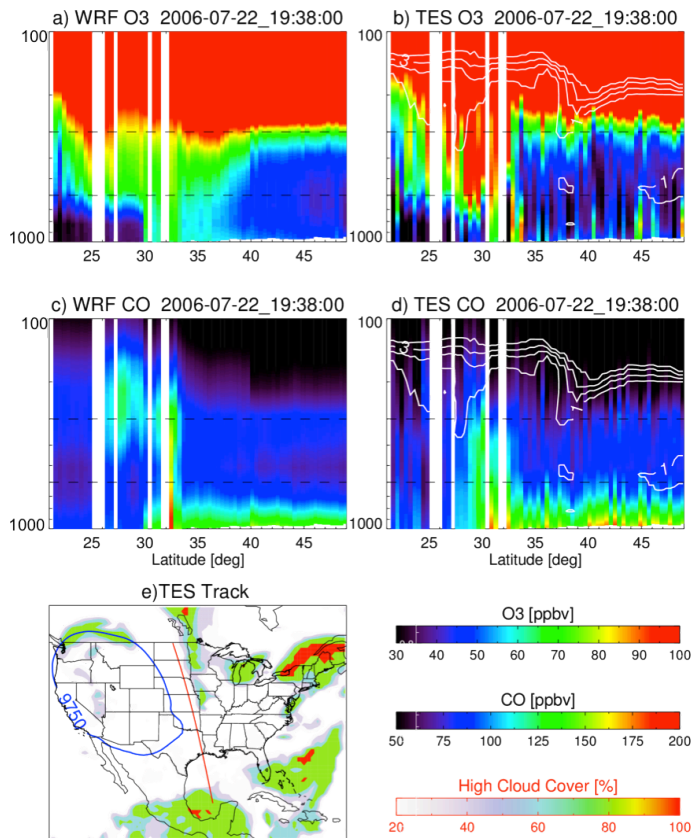
**Fig. 4.** Frequency distribution of radar reflectivity from observations (black lines) and model results (blue lines) at different hours of the day (see legend; time is UTC) for the 15 July to 6 August period. **(a)** Rocky Mountain region, **(b)** Gulf Coast region, **(c)** Central Plains region, and **(d)** Midwest region as shown in Fig. 1.

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**Fig. 5.** Modeled  $\text{O}_3$  (left panels) and CO (right panels) at 300 hPa averaged for the 4 time periods discussed in the text. The 9725 and 9750 m heights indicate the location of the upper troposphere anticyclone. The black patch in Northwest Mexico (panels **g** and **h**) is a region with geopotential height of 9725 m.



**Fig. 6.** O<sub>3</sub> (panels **a** and **b**) and CO (panels **c** and **d**) from the WRF-Chem simulation (panels **a** and **c**) and the AURA/TES instrument (panels **b** and **d**). The horizontal dashed lines mark the 300 and 600 hPa levels. Results are plotted along the TES overpass track for 22 July 2006 over the central US as indicated in panel (**e**). The potential vorticity (panels **b** and **d**; 1, 2, 3, 4 PVU contours plotted), geopotential height, and high cloud cover (panel **e**) from the ERA-Interim 00:00 UTC analysis are also shown.

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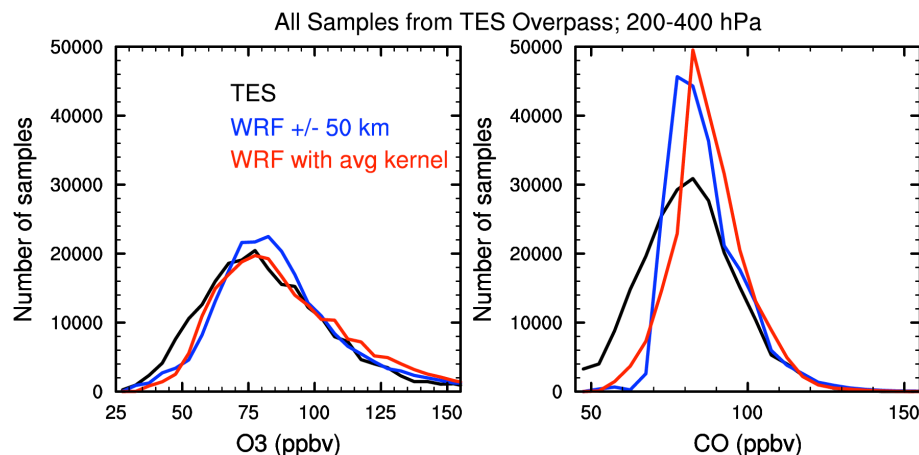
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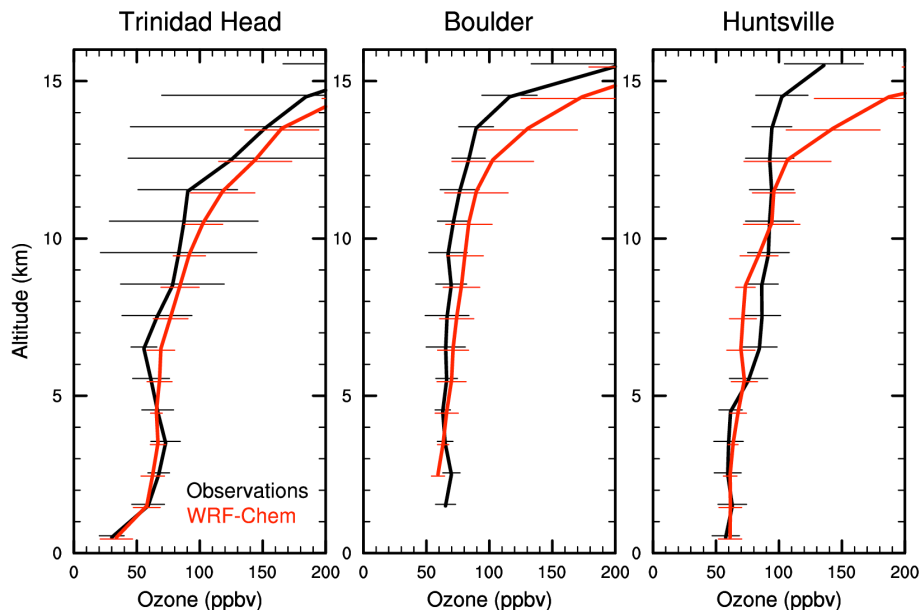
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**Fig. 7.** Frequency distribution of O<sub>3</sub> (left panel) and CO (right panel) for the 200–400 hPa altitudes from TES data (black line), WRF-chem results (blue line), and WRF-chem modified by the TES a priori and averaging kernel (red line) for 15 July to 6 August 2006 TES passes over the model domain.

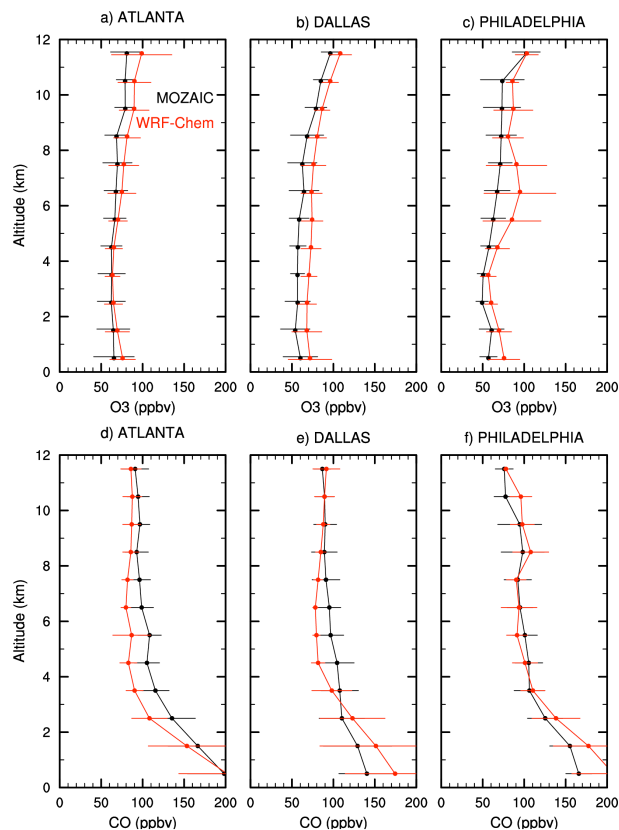
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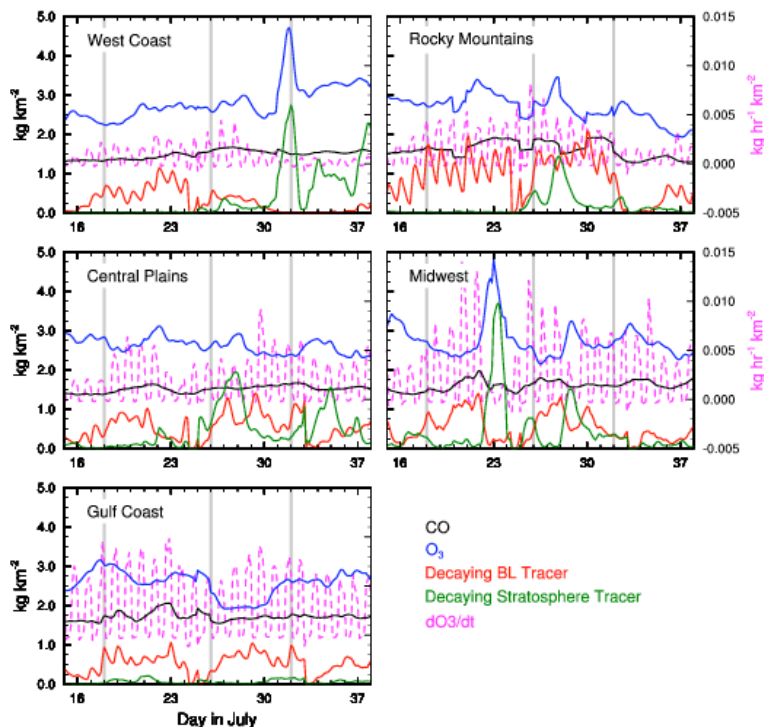
**Fig. 8.** Average  $\text{O}_3$  profiles from ozonesondes (black lines) and WRF-Chem (red lines) for (a) Trinidad Head, California, (b) Boulder, Colorado, and (c) Huntsville, Alabama. The profiles are averages over the 15 July to 6 August 2006 when ozonesondes were launched. The standard deviations for each of the profiles are indicated by the horizontal lines.

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**Fig. 9.** Average vertical profiles of  $O_3$  (a–c) and CO (d–f) mixing ratios from the MOZAIC data (black lines) and the WRF-Chem simulation (red lines) binned into altitude levels. The profiles are averages over the 15 July to 3 August 2006 when MOZAIC profiles were available. The standard deviation for each of the datasets is indicated by the horizontal lines.

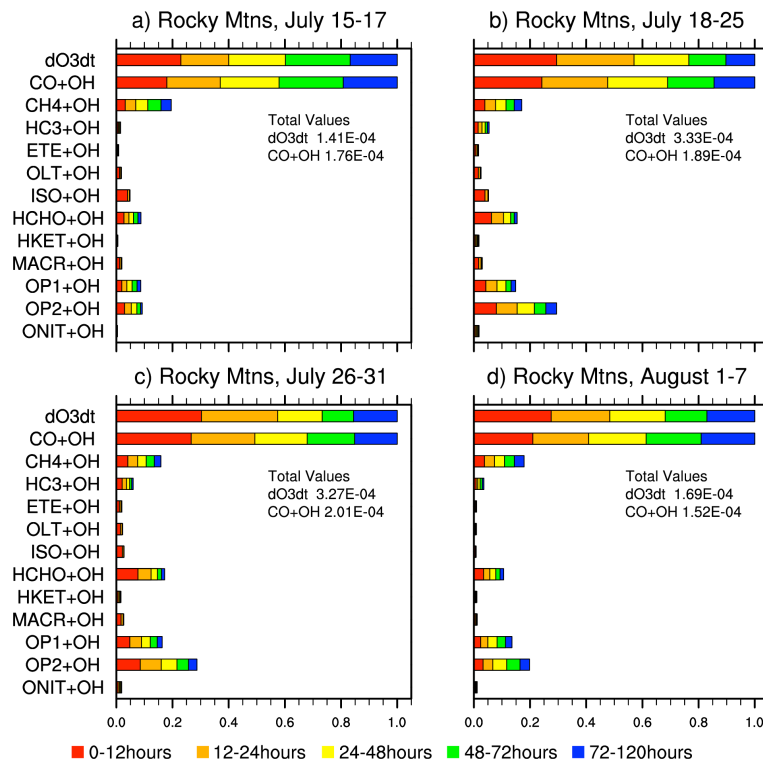




**Fig. 10.** Mass per unit area of CO (black line), O<sub>3</sub> (blue line), decaying boundary layer tracer (red line), decaying stratosphere tracer (green line), and chemical ozone tendency (dashed magenta line; right axis) in the upper troposphere (8–12 km altitude range) over the (a) West Coast, (b) Rocky Mountains, (c) Central Plains, (d) Midwest, and (e) Gulf Coast regions of the model domain as shown in Fig. 1. The decaying boundary layer scalar and decaying stratosphere tracer masses per area are the respective masses of air in units of  $2 \times 10^6 \text{ kg km}^{-2}$  and  $1 \times 10^6 \text{ kg km}^{-2}$ , respectively. The vertical gray lines delineate the 4 meteorological periods as described in the text.

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**Fig. 11.** The contribution of various VOC oxidation reactions to peroxy radical formation is shown as a function of the age of air for the Rocky Mountain region. Each contribution is an average for the upper troposphere (8–12 km altitude range) over the four time periods described in the meteorology section of the paper. The  $dO_3dt$  and  $CO + OH$  rates are normalized by their total value (listed in the panel) in the region and time period, while the other reaction rates are normalized by the  $CO + OH$  total value in the region and time period. The Rocky Mountain region is shown in Fig. 1.

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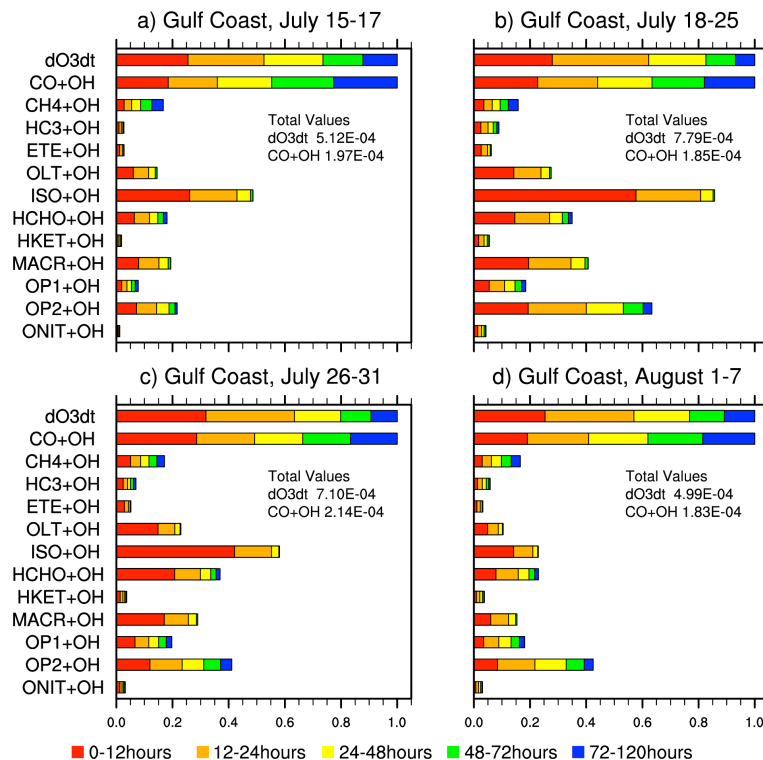
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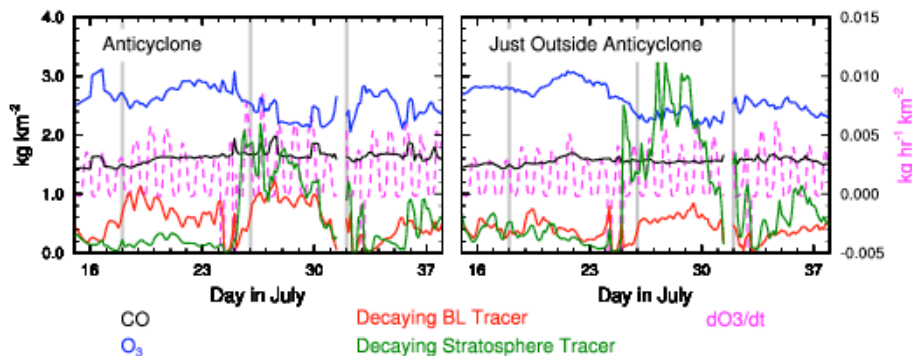
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**Fig. 12.** The same as Fig. 11 but for the Gulf Coast region, which is shown in Fig. 1.

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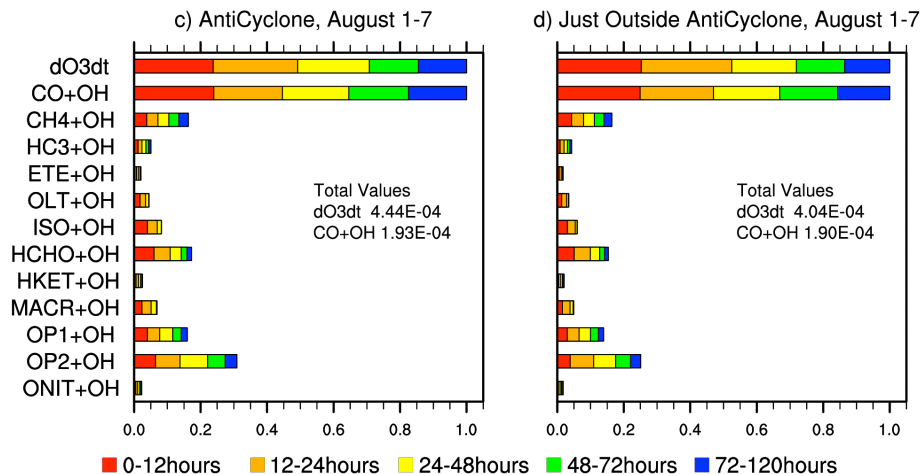


**Fig. 13.** Mass per unit area of CO (black line), O<sub>3</sub> (blue line), decaying boundary layer tracer (red line), and decaying stratosphere tracer (green line), and chemical ozone tendency (dashed magenta line; right axis) in the upper troposphere (8–12 km altitude range) within the anticyclone and just outside the anticyclone. The horizontal area of the anticyclone is defined as the region by  $z > 9730$  m at 300 hPa. The horizontal area of the just outside anticyclone region is defined as the region by  $9730 \text{ m} > z > 9680$  m at 300 hPa and only includes grid points within 600 km of the anticyclone. The boundary layer scalar and stratosphere tracer masses per area are the respective masses of air in units of  $2 \times 10^6 \text{ kg km}^{-2}$  and  $1 \times 10^6 \text{ kg km}^{-2}$ , respectively. On the 31 July, geopotential heights were  $< 9730$  m, so the anticyclone could not be located.

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**Fig. 14.** The contribution of various VOC oxidation reactions to peroxy radical formation is shown as a function of the age of air for inside (left panels) and just outside (right panels) the UT anticyclone. Each contribution is averaged in the upper troposphere (8–12 km altitude range) over the one of the time periods described in the meteorology section of the paper. The  $dO_3dt$  and  $CO + OH$  rates are normalized to their total value (listed in the panel) in the region and time period, while the other reaction rates are normalized to the  $CO + OH$  total value in the region and time period. The horizontal area of the anticyclone is defined as the region by  $z > 9730$  m at 300 hPa. The horizontal area of the just outside anticyclone region is defined as the region by  $9730 \text{ m} > z > 9680$  m at 300 hPa and only includes grid points within 600 km of the anticyclone.

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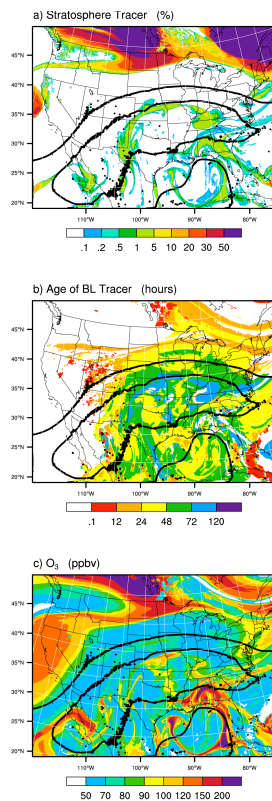
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**Fig. 15.** WRF-Chem simulated **(a)** decaying stratosphere tracer (%), **(b)** age of the boundary layer tracer (hours), and **(c)** O<sub>3</sub> mixing ratio (ppbv) at 250 hPa for 21:00 UTC 5 August 2006. Overlaid is geopotential height at 300 hPa for  $z = 9680$  m and  $9730$  m, which defines the anticyclone in the analysis.

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