



Peroxy Radical Measurements by Ethane - Nitric Oxide Chemical Amplification and Laser-Induced Fluorescence / Fluorescence Assay by Gas Expansion during the IRRONIC field campaign in a Forest in Indiana

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Abstract. Peroxy radicals were measured in a mixed deciduous forest atmosphere in Bloomington, Indiana, USA, during the Indiana Radical, Reactivity and Ozone Production Intercomparison (IRRONIC) during the summer of 2015. Total peroxy radicals ($[XO_2] \equiv [HO_2] + \Sigma[RO_2]$) were measured by a newly developed technique involving nitric oxide (NO) – ethane (C_2H_6) chemical amplification followed by NO_2 detection by cavity attenuated phase shift spectroscopy (hereinafter referred to as ECHAMP). The sum of hydroperoxy radicals (HO_2) and a portion of organic peroxy radicals ($[HO_2^*] = [HO_2] + \Sigma\alpha_i[R_iO_2]$, $0 < \alpha < 1$) was measured by the Indiana University Laser-Induced Fluorescence / Fluorescence Assay by Gas Expansion instrument (LIF-FAGE). Additional collocated measurements include concentrations of NO, NO_2 , O_3 , and a wide range of volatile organic compounds (VOCs); and meteorological parameters. XO_2 concentrations measured by ECHAMP peaked between 13:00 to 16:00 local time, with campaign average concentrations of 41 ± 15 ppt (1σ) at 14:00. Daytime concentrations of isoprene averaged 3.6 ± 1.9 ppb (1σ) whereas average concentrations of NO_x ($[NO] + [NO_2]$) and toluene were 1.2 ppb and 0.1 ppb, respectively, indicating a low impact from anthropogenic emissions at this site.

We compared ambient measurements from both instruments and conducted a calibration source comparison. For the calibration comparison, the ECHAMP instrument, which is primarily calibrated with an acetone photolysis method, sampled the output of the LIF-FAGE calibration source which is based on the water vapor photolysis method and, for these comparisons, generated a 50-50% mixture of HO_2 and either butane or isoprene-derived RO_2 . A bivariate fit of the data yields the relation $[XO_2]_{ECHAMP} = 0.88 ([HO_2] + [RO_2])_{IU_cal} + 6.6$ ppt, with an R^2 of 0.99. This level of agreement is within the combined analytical uncertainties for the two instruments' calibration methods.

A linear fit of all the 30-minute averaged $[XO_2]$ ambient data with the 1-minute averaged $[HO_2^*]$ data (one point per 30 minutes) yields the relation $[XO_2] = 0.85 [HO_2^*] + 3.3$ ppt, with an R^2 of 0.67. Day to day variability in the $[XO_2]/[HO_2^*]$ ratio was observed. The lowest $[XO_2]/[HO_2^*]$ ratios between 13:00 and 16:00 were 0.8 on 13 and 18 July, whereas the highest ratios of 1.1 to 1.3 were observed on 24 and 25 July – the same two days on which the highest concentrations of isoprene and ozone were observed. Although the exact composition of the peroxy radicals during IRRONIC is not known, 0-dimensional photochemical modeling of the IRRONIC dataset using the RACM2, RACM2-LIM1, MCM 3.2, and MCM 3.3.1 chemical mechanisms predicts afternoon $[XO_2]/[HO_2^*]$ ratios of between 1.1 to 1.5 depending on mechanism. Differences between the observed ambient $[XO_2]/[HO_2^*]$ ratio and that predicted with the 0-D modeling can be attributed to deficiencies in the model, errors in the two measurement techniques, or both. Although these comparison results are encouraging and demonstrate the viability of using the new ECHAMP technique for field measurements of peroxy radicals, further research investigating the overall accuracy of the measurements and possible interferences from both methods is warranted.

30 1. Introduction

Peroxy radicals in the atmosphere comprise hydroperoxy (HO_2) and organic peroxy radicals (RO_2 , R = organic group). The most important sources of peroxy radicals are the reactions of oxidants (OH, O_3 , and NO_3) with volatile organic



compounds (VOCs), photolysis of oxygenated VOCs, and decomposition of peroxyacetyl nitrate (PAN) (Atkinson, 2000). Chemistry involving “RO_x” radicals ($[RO_x] \equiv [OH] + [RO_2] + [HO_2]$) leads to the formation of ozone (O₃), oxygenated VOCs, and secondary aerosol particles (Kroll and Seinfeld, 2008; Ng et al., 2008; Atkinson and Arey, 2003; Akimoto, 2016; Claeys et al., 2004). The chemical identity and concentrations of peroxy radicals can provide important information on atmospheric oxidation processes such as ozone production, the removal efficiency of primary pollutants, and radical budgets. This information is ultimately required to formulate pollution control strategies and to evaluate the impacts of atmospheric chemistry on health and global climate. It is therefore crucial to understand the concentrations and chemistry of RO_x radicals in the atmosphere.

Comparison of measured radical concentrations to those produced by photochemical models is a common exercise used to assess our understanding of atmospheric chemistry. Discrepancies of a factor of two or more between measured and modeled OH concentrations have been reported in biogenic VOC-rich forest environments (Lelieveld et al., 2008; Lu et al., 2012; Pugh et al., 2010), suggesting that our knowledge of atmospheric photochemistry is deficient. Similarly, discrepancies between measured and modeled peroxy radicals have suggested the presence of unknown sources or sinks of peroxy radicals (Griffith et al., 2013; Wolfe et al., 2014). These findings have fueled research into the oxidation mechanisms of biogenic VOCs, especially isoprene (e.g., Wennberg et al., 2018). Although much has been learned in the past decade, the atmospheric fate of biogenic VOCs remains incompletely understood.

Some past model-measurement comparisons are difficult to interpret because of measurement errors that have recently been discovered. Measurements of OH by the laser-induced fluorescence technique can be affected by a sampling-related interference which can exceed the actual concentration of OH (Mao et al., 2012). Similarly, many previous measurements of HO₂ by chemical conversion to OH through the HO₂ + NO → OH + NO₂ reaction using both the LIF-FAGE and the perCIMS techniques have been shown to have been affected by a variable contribution from organic peroxy radicals (Fuchs et al., 2011; Hornbrook et al., 2011) and the LIF-based measurements subject to this interference are now referred to as HO₂* ($[HO_2^*] \equiv [HO_2] + \alpha_i \sum [R_iO_2]$, $0 < \alpha < 1$). The sensitivity of the LIF-FAGE technique to each type of organic peroxy radical varies with the amount of NO added for the conversion and is instrument-dependent but in general is highest (~90%) for β-hydroxy peroxy radicals derived from alkenes and lowest for those derived from small alkanes (Fuchs et al., 2011; Lew et al., 2018; Whalley et al., 2013).

Discrepancies between measured and model-predicted OH and XO₂ concentrations can be caused by a combination of measurement errors, missing or incorrect chemistry in models and erroneous model constraints. Measurement errors can be evaluated by the comparison of atmospheric measurements by multiple techniques. Several HO_x intercomparison projects have been conducted in the past few decades (Eisele et al., 2003; Fuchs et al., 2010; Hofzumahaus et al., 1998; Ren et al., 2012; Schlosser et al., 2009). There have been few intercomparisons, however, of total peroxy radical ($[HO_2] + \sum [RO_2]$) measurements and these have produced mixed results. For example, excellent agreement between the matrix isolation electron spin resonance (MI-ESR) and the RO_x LIF-FAGE techniques was observed in a chamber study involving HO₂, CH₃O₂, and C₄H₇O₂ produced by the oxidation of methane and 1-butene (Fuchs et al., 2009). An earlier comparison of XO₂



measurements between a CO-based chemical amplifier (PERCA) and MI-ESR showed overall agreement of within 10% (Platt et al., 2002). In contrast, XO₂ measurements in a forest from two similar CO-based chemical amplifiers differed by more than a factor of three (Burkert et al., 2001). This disagreement was attributed to sampling losses on a rain cover. Similarly, XO₂ measurements from two CO-based chemical amplifiers during the airborne African Monsoon Multidisciplinary Analysis (AMMA) campaign differed by factors of 2-4 when the usual relative humidity-dependent calibration (Mihele and Hastie, 1998) was used for the chemical amplifier data (Andrés-Hernández et al., 2010). As a result, the relative humidity dependence of the chemical amplification technique has been questioned (Andrés-Hernández et al., 2010; Sommariva et al., 2011) despite strong experimental evidence (Butkovskaya et al., 2007; Butkovskaya et al., 2005; Butkovskaya et al., 2009; Mihele et al., 1999; Mihele and Hastie, 1998). Due to the paucity of XO₂ measurement intercomparisons and these new questions regarding the impact of relative humidity on the traditional chemical amplifier technique, further intercomparisons involving different instruments are required before we have enough confidence in the measurements to interpret model-measurement discrepancies as arising from unknown chemistry in models.

This paper presents measurements of XO₂ in a mixed deciduous forest by the new Ethane CHEMical AMPlifier (ECHAMP) technique (Wood et al., 2017), which is a variation of the traditional chemical amplification or “PERCA” method (Cantrell and Stedman, 1982; Hastie et al., 1991; Wood and Charest, 2014). Measured XO₂ concentrations at this high isoprene, low NO_x environment are described along with supporting measurements of ozone (O₃), nitrogen oxides (NO_x), biogenic and anthropogenic VOCs, and meteorology. We compare measurements of XO₂ by ECHAMP with collocated ambient measurements of HO₂* by the Indiana University LIF-FAGE technique. We also describe calibration comparison experiments in which ECHAMP, which was calibrated by an acetone photolysis calibration method, quantified radical concentrations produced by the LIF-FAGE calibration source which is based on the more common water photolysis method. Ozone formation rates are also calculated based on measured XO₂ and NO concentrations.

2 Experimental Section

2.1 Site description

The measurements were carried out at the Indiana University Research and Teaching Preserve (IURTP) field laboratory during the Indiana Radical, Reactivity and Ozone Production Intercomparison (IRRONIC) campaign over the time period of 9 July – 8 August 2015. The IURTP is located in a mixed deciduous forest 1 km from the perimeter road for Indiana University in Bloomington (Fig. 1). Instrument inlets and related instrumental accessories were set atop a 3 meter scaffolding platform in a clearing behind the IURTP building. The height of the scaffolding was several meters below the forest canopy. The major analytical instruments and gas cylinders were housed inside the building.



2.2 ECHAMP Measurements of Total Peroxy Radicals (XO₂)

XO₂ concentrations were quantified using a newly developed analytical technique, which involves chemical amplification by ethane (C₂H₆) - nitric oxide (NO) followed by nitrogen dioxide (NO₂) detection using cavity attenuated phase shift spectroscopy (hereinafter referred as ECHAMP: Ethane CHEMICAL AMPLIFIER) (Wood et al., 2017). This instrument can be thought of as a descendent of “traditional” chemical amplifiers, also known as PERCA, in which ambient air is mixed with carbon monoxide and nitric oxide and the resulting “amplified” NO₂ measured by the luminol technique (Cantrell and Stedman, 1982; Clemitshaw et al., 1997; Kartal et al., 2010; Mihele and Hastie, 2000). Our initial peroxy radical sensor (Wood and Charest, 2014) relied on the original CO/NO amplification chemistry but utilized a modern, highly sensitive NO₂ detection method: cavity attenuated phase shift spectroscopy (CAPS) (Kebabian et al., 2007; Kebabian et al., 2008). The major modification made for the ECHAMP method used for the measurements described in this study is that ethane (C₂H₆) replaces CO as a reagent. This results in greatly improved deployability thanks to the relative safety of C₂H₆ compared to CO, a smaller dependence of the sensitivity on relative humidity, but at the expense of lower amplification factors. Details of the experimental technique are described elsewhere (Wood et al., 2017) but its deployment at the IURTP is described here.

The ECHAMP inlet was attached to scaffolding at a height of 3 m. Ambient air was sampled at a flow rate of 5.5 standard liters per minute (SLPM) into a 0.4 cm inner diameter (ID) glass sampling cross internally coated with halocarbon wax (Halocarbon Products Corp., series 1500) and externally coated with PTFE tape. The sampled air then entered two identical reaction chambers (0.4 cm ID × 61 cm, FEP tubing) at a flowrate of 0.87 SLPM - see schematic in Wood and Charest (2014). The total residence time in the sampling cross before entering the reaction chambers was approximately 18 ms.

At any given point in time, one reaction chamber operated in “amplification” (RO_x) mode while the other operated in “background” (O_x) mode. In “RO_x” mode, the air was immediately mixed with NO and C₂H₆ in the “upstream” reagent addition port and, 0.1 second later, mixed with nitrogen (N₂) in the “downstream” reagent addition port, effecting the following radical propagation reactions:



Reactions R3 through R7 repeat several times, leading to the formation of NO₂ that is subsequently measured by a CAPS sensor. In background (O_x) mode, the N₂ and C₂H₆ flows were switched: sampled air was mixed with NO and N₂ upstream and C₂H₆ downstream. During this sampling mode, sampled radicals are removed by a combination of reactions R1, R2, R3 and finally the reaction of OH with NO to form HONO. The flowrates of NO, N₂ and C₂H₆ were each maintained at 45 sccm



using mass flow controllers (MKS model 1179A and Alicat MC series). Cylinder concentrations of NO and C₂H₆ (Indiana Oxygen) were 21.1 ppm and 30%, respectively, leading to concentrations in the reaction chamber of 0.9 ppm and 1.4%, respectively. Both upstream and downstream injections were delivered with PFA tubing (0.16 cm i.d. × 6 m). Each reaction chamber alternated between RO_x and O_x mode every 45 seconds on an anti-synchronized schedule using four solenoid valves controlled by Labview software (National Instruments). After the downstream reagent addition, the air from each reaction chamber flowed through 1 m of 0.32 cm ID FEP tubing, a particulate matter filter (United Filtration Systems, Inc., DIF BN60), and another 6 m of tubing before entering identical CAPS monitors located inside the laboratory. The CAPS NO₂ measurements during “RO_x” mode are from ambient NO₂, NO₂ from the reaction of NO and O₃ in the reaction chamber and transport tubing, and NO₂ from the chemical amplification reactions involving HO₂ and RO₂ (R1 through R7). In “O_x” mode, the CAPS measures NO₂ from the first two categories above and NO₂ produced by R1 and R3 but not from the amplification reactions (R3 to R7), as ethane is not added until all radicals are removed by formation of HONO. The concentrations of peroxy radicals were calculated by dividing the difference between the two CAPS sensors’ NO₂ measurements (ΔNO₂) between “RO_x” and “O_x” modes by an experimentally determined amplification factor F:

$$15 \quad [\text{RO}_2] + [\text{HO}_2] = \Delta[\text{NO}_2]_{(\text{CAPS A} - \text{CAPS B})}/F \quad (1)$$

The RH-dependent amplification factor F was measured using the acetone photolysis method described by Wood and Charest (2014). Briefly, methyl peroxy (CH₃O₂) and peroxyacetyl (CH₃C(O)OO) radicals (50 – 400 ppt) were produced by the photolysis of acetone vapor and reacted with excess NO to form NO₂ which was quantified using a CAPS NO₂ sensor. The accuracy of this calibration method ultimately depends on the accuracy of the CAPS NO₂ measurement (see supplementary information (SI)) and knowledge of the products of the reaction of CH₃O₂ and CH₃C(O)OO with NO but does not depend on measurements of actinic flux.

The amplification factor F was measured to be 28 at 0% relative humidity (RH) and decreased to 6 at 90% RH (Wood et al., 2017). The RH was typically between 50 and 75% during the afternoon, corresponding to values of F between 20 and 11. These values are based on laboratory calibrations performed before and after the field project. During the field campaign, we attempted to use a variation on the calibration method described by Wood and Charest (2014). Rather than flow air through the headspace over pure acetone to produce dilute acetone vapor, we instead flowed air through the headspace of dilute (1%) aqueous acetone in an attempt to obviate the need to dilute the resulting acetone vapor (i.e., by reducing the vapor pressure of the acetone per Raoult’s Law). Inconsistent calibrations resulted, however, and subsequent laboratory tests demonstrated that the use of aqueous acetone sometimes produced compounds that absorb blue light and therefore interfered with the CAPS NO₂ measurement which is based on absorption of light at 450 nm with a bandpass of 10 nm (full width at half maximum). Because field calibrations were unsuccessful, we have increased the measurement uncertainty accordingly (see below). The acetone vapor photolysis calibration results obtained in the laboratory also agreed



with our prototype H₂O photolysis method as described in Wood et al (2017). Further details on the calibration are described in the SI.

Individual peroxy radicals are not detected with equal sensitivity by ECHAMP due to the formation of organic nitrates and organic nitrites in the reaction chambers:

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Including a sampling loss term, the sensitivity “ α ” of ECHAMP to individual organic peroxy radicals relative to that of HO₂ can be summarized by Equation 2:

$$\alpha_{\text{RiO}_2} = S_{\text{RiO}_2}/S_{\text{HO}_2} = L_i(1-Y_i)(k_{\text{Ri9a}}[\text{O}_2]/(k_{\text{Ri9a}}[\text{O}_2] + k_{\text{Ri9b}}[\text{NO}])) \quad (2)$$

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where $S_{\text{RiO}_2}/S_{\text{HO}_2}$ is the sensitivity of ECHAMP to individual RO₂ compounds relative to that of HO₂, L_i is the fractional sampling transmission of an individual organic peroxy species “R_iO₂” through the short inlet into the reaction chambers (relative to that of HO₂), Y is the alkyl nitrate yield ($Y = \text{R8b}/(\text{R8a} + \text{R8b})$), and the remaining terms in parentheses account for alkyl nitrite (RONO) formation. Alkyl nitrate yields increase with carbon backbone number, from less than 0.1% for CH₃O₂ to 8% for isoprene to over 25% for C10 and larger alkyl peroxy radicals (Lockwood et al., 2010; Orlando and Tyndall, 2012). Alkyl nitrite (RONO) formation accounts for less than 4% loss for most organic peroxy radicals and is likely negligible for alkene-derived peroxy radicals due to the rapid decomposition of beta hydroxy alkoxy radicals (Atkinson, 1997), but can sequester a calculated 10% of CH₃O₂ (Wood et al., 2017). Sampling losses are limited to the 18 ms transit time in the halocarbon wax-coated sampling cross to the tee in which the NO and C₂H₆ are added. Mihele et al. (1999) measured effective first order wall loss rate constants of 3 to 7 s⁻¹ for HO₂ onto ¼” OD PFA tubing, depending on RH, and ~0.5 s⁻¹ for CH₃O₂ and C₂H₅O₂. Though this would suggest losses in our inlet of up to 12% for HO₂ and 1% for the alkyl peroxy radicals, laboratory tests on our inlet have demonstrated losses of less than 2% for HO₂ in our inlet and loss rate constants onto various fluoropolymers much lower than presented in Mihele et al. (1999) as described in the SI.

At an RH of 50%, the theoretical 1 σ precision of the ECHAMP measurements, limited by only the precision of the CAPS NO₂ measurements and the amplification factor, was 0.8 ppt for a 90-second average. The atmospheric variability of O₃, which after reaction with NO accounts for most of the NO₂ observed by the CAPS sensors, led to an additional contribution to the noise due to the slightly different time responses of the two CAPS sensors. The observed precision during sampling was typically 2.5 ppt (1 σ) for 90-second averaging (Wood et al., 2017), leading to a detection limit of 5 ppt for 90-second averaging and 1.6 ppt for 15 minute averages at a signal-to-noise ratio of two. At night, although variability of O₃



was negligible, high RH values of over 95% and the resulting low values of F led to detection limits of between 2 ppt and 8 ppt for 90 second average measurements.

We assign an uncertainty of 27% (2σ) to the ECHAMP measurements during the IRRONIC project, comprising the uncertainty in the NO_2 calibration of the CAPS sensors (5%), the uncertainty in the relative humidity - dependent amplification factor (usually 16%, but increased to 25% because post-deployment laboratory calibrations were used instead of the unsuccessful field calibrations using aqueous acetone), and the variable sensitivity to speciated peroxy radicals. These uncertainties are more fully described in Wood et al. (2017). Except where noted otherwise, all ECHAMP XO_2 measurements presented are 15-minute averages.

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2.3 Laser-Induced Fluorescence Measurements HO_2^*

HO_2^* was measured by the Laser-Induced Fluorescence / Fluorescence Assay by Gas Expansion (LIF-FAGE) technique described in detail elsewhere (Griffith et al., 2013a; Dusanter et al., 2009). Briefly, air is sampled through a pinhole into a low pressure chamber and mixed with NO which converts HO_2 into OH. OH radicals are excited by 308 nm radiation from a tunable dye laser and the subsequent fluorescence detected with a time-gated microchannel plate photomultiplier (MCP-PMT) detector. Some organic peroxy radicals are also converted into OH in the LIF-FAGE instrument. Based on laboratory tests, the sensitivities “ α ” of the LIF-FAGE measurement for the added NO concentrations used in this study relative to HO_2 for the following RO_2 radicals are 83% for isoprene- RO_2 , 91% for methyl vinyl ketone RO_2 , 54% for methacrolein RO_2 , 65% for ethene- RO_2 , 65% for toluene- RO_2 , 15% for propane- RO_2 , and 31% for butane- RO_2 (Lew et al., 2018). The conversion efficiency for other major RO_2 can be either estimated or measured to be 5% for CH_3O_2 and the acetyl peroxy radical ($\text{CH}_3\text{C}(\text{O})\text{O}_2$), 8% for ethyl peroxy radical ($\text{C}_2\text{H}_5\text{O}_2$), and 31-55% for RO_2 compounds from the OH oxidation of high-molecular-weight hydrocarbons (Fuchs et al., 2011; Griffith et al., 2013). These conversion efficiencies are average values weighted over the distribution of isomers where applicable.

The LIF-FAGE instrument was calibrated using a portable calibrator in which quantified amounts of OH/ HO_2 and RO_2 were produced through the photolysis of water vapor by a low-pressure mercury lamp at 184.9 nm (Dusanter et al., 2008). Humid air containing either isoprene (80 ppb) or n-butane (1.4 ppm) entered the rectangular calibrator ($1.27 \times 1.27 \times 30$ cm). Light from a low-pressure mercury lamp (UVP Inc, model 11sc1) illuminated a ~ 3 cm³ photolysis volume through a quartz window. The flow rate of air was maintained at 45 SLPM. A mixture with equal concentrations of HO_2 and either $\text{C}_5\text{H}_8(\text{OH})\text{O}_2$ (from isoprene) or $\text{C}_4\text{H}_9\text{O}_2$ (from butane) were produced when isoprene or butane were added to the calibration gas upstream of the photolysis region, respectively. Ozone actinometry was used to quantify the product of the actinic flux and the exposure time (“Ft”) in the calibrator (Dusanter et al., 2008). Concentrations of generated peroxy radicals are calculated by the following equation :



$$[HO_2] + [RO_2] = \frac{[O_3][H_2O]\sigma_{H_2O}\phi_{H_2O}}{[O_2]\sigma_{O_2}\phi_{O_2}} \quad (3)$$

where $[O_3]$ is the concentration of ozone generated by the photolysis of O_2 ; σ_{H_2O} and σ_{O_2} are the absorption cross sections of H_2O and O_2 at 184.9 nm, respectively; and ϕ_{H_2O} and ϕ_{O_2} are the photolysis quantum yields, both equal to two (Washida et al., 1971). A value of $7.14 \times 10^{-20} \text{ cm}^2 \text{ molecule}^{-1}$ (base e) was used for σ_{H_2O} (Cantrell et al., 1997; Hofzumahaus et al., 1997). The effective value of σ_{O_2} depends on the O_2 optical depth and the operating conditions of the mercury lamp and was determined to be $1.20 \times 10^{-20} \text{ cm}^2 \text{ molecule}^{-1}$ (Dusanter et al., 2008; Lanzendorf et al., 1997). The water vapor mixing ratio was measured by IR absorption spectrometry using a LI-COR 6262 monitor. Ordinarily the ozone mixing ratio is determined using a calibrated photodiode installed in the calibrator (Griffith et al., 2013). The conversion factor (calibration) that converts the photodiode reading to an O_3 mixing ratio is determined from separate experiments in which a range of O_3 concentrations produced by the calibrator are measured with a UV-absorption O_3 sensor. For this project, $[O_3]$ was instead quantified by the ECHAMP CAPS NO_2 sensors after conversion to NO_2 by reaction with excess NO. This produces a very precise measurement of the sum of $[O_3]$ and $[NO_2]$ (1σ precision of 22 ppt for 10 second averages). The accuracy of this ozone determination is thus ultimately traceable to the CAPS NO_2 calibration (see SI). Linking the IU FAGE HO_2^* calibration to the ECHAMP NO_2 measurement has ramifications for the intercomparison of the IU calibration source and the ambient measurements as discussed in the relevant sections below.

The sensitivity of the instrument is corrected for fluorescence quenching by water vapor. This amounted to a correction of approximately 20% at a water mixing ratio of 1%. The limit of detection of HO_2^* was 0.8 ppt (30 s average, signal-to-noise ratio of two). The overall accuracy of the HO_2^* measurements was $\pm 36\%$ (2σ). On all days except 22 July, HO_2^* data were collected for 1 minute every 30 minutes and OH was measured during the rest of the 30 minute cycle. On 22 July, OH was not measured and instead the FAGE instrument measured HO_2^* continuously.

2.4 Supporting Measurements

Ambient NO_2 was measured using a separate CAPS monitor (Aerodyne Research) (Kebabian et al., 2007; Kebabian et al., 2008). The standard 450 nm bandpass filter used by the CAPS monitor was replaced with a 470 nm bandpass filter to eliminate any interference by glyoxal and methyl glyoxal (Kebabian et al., 2008). This reduced the sensitivity by approximately a factor of three but still provided high signal-to-noise ratios (>100) for the ambient measurements. O_3 was measured with a UV absorbance monitor (2B Technologies model 202). NO was measured using a Thermo Fisher chemiluminescence sensor (Model 42i Trace Level). NO, NO_2 , and O_3 data were averaged to 1 minute. Additional details regarding the calibrations and baseline measurements for the NO, NO_2 , and O_3 measurements can be found in the SI.

A wide variety of biogenic and anthropogenic VOCs including isoprene and its oxidation products (methyl vinyl ketone and methacrolein), monoterpenes, non-methane hydrocarbons (C_2 - C_5 and C_6 - C_{12}), including aromatics, and



oxygenated VOCs (alcohols, aldehydes and ketones) were measured during IRRONIC. An online GC-FID-FID was used to measure 57 NMHCs ([Badol et al., 2004](#)). Ambient air was sampled through a NAFION membrane and NMHCs were trapped at a temperature of -30 °C inside a quartz tube filled with Carbosieve SIII and Carbopack B. A thermodesorption unit (Perkin Elmer, ATD 400) was used to inject the sample into two columns (PLOT alumine and CPSil 5CB) to separate C₂-C₆ and C₆-C₁₂ compounds. Two FID detections provided limits of detection of 10–60 pptv at a time resolution of 90 min. A second online GC-FID instrument was used to measure ethanol, isopropanol, methylethylketone and a few monoterpenes (α -pinene, 3-carene) ([Roukos et al., 2009](#)). A sampler unit (Markes International, air server Unity 1) allowed continuous sampling of ambient air through a trap held at 12 °C and filled with Carbopack B and Carbopack X. After thermodesorption, the GC separation was performed using a high-polarity CP-Lowox column (Varian, France). Limits of detection reached with this instrument were in the range 10–90 pptv for a time resolution of 90 min. Offline sampling was performed on multisorbent cartridges to measure > C₉ anthropogenic compounds (alkanes and aromatics) and monoterpenes (pinenes, terpinenes, limonene, ocimene, terpinolene, camphene, myrcene, borneol, camphor, cumene), and on DNPH (DiNitroPhenylHydrazine) cartridges to measure carbonyls, including formaldehyde, acetaldehyde and higher compounds. The cartridge measurements were integrated over 2-h sampling periods. Technical details can be found in ([Ait-Helal et al., 2014](#); [Detournay et al., 2011](#); [Detournay et al., 2013](#)).

Zero-dimensional photochemical modeling of this field campaign data was performed using the Framework for 0-Dimensional Atmospheric Modeling (FOAM) which was constrained by the 30 minute average mixing ratios of the supporting measurements ([Wolfe et al., 2016](#)). FOAM was executed using four different chemical mechanisms: two versions of the Regional Atmospheric Chemistry Mechanism (RACM2 and RACM2-LIM1) and the Master Chemical Mechanism (MCM 3.2 and 3.3.1). RACM2 groups various compounds based on similar rates of reaction resulting in 363 reactions from 17 stable inorganics, 4 inorganic intermediates, 55 stable organics, and 43 intermediate organics ([Goliff et al., 2013](#)). RACM2-LIM1 incorporates the revision to the isoprene oxidation mechanism ([Peeters et al., 2009](#)) that includes the Leuven Isoprene Mechanism (LIM) including a 1,6 H-shift and a 1,5 H-shift for isoprene peroxy radicals. MCM is a near-explicit chemical reaction model resulting in approximately 17000 reactions from 6700 radical species from methane and 142 non-methane species. Similar to the LIM1 mechanism, MCM 3.3.1 was updated to include revisions to the isoprene oxidation mechanism resulting in HO_x recycling from peroxy radical H-shift isomerization as well as NO_x recycling and updated ozonolysis rate constants.

3 Results and Discussion

3.1 Calibration comparisons between ECHAMP and IU calibration source

On 24 and 26 July the IU calibration source was positioned so that its output overflowed the ECHAMP inlet. Figure 2 compares the response of ECHAMP to variable concentrations of peroxy radicals generated by the IU calibrator. Concentrations of peroxy radicals were varied by adjusting the mixing ratio of water or by changing the intensity of the UV



lamp. H₂O mixing ratios varied from 0.1 to 1.4%, corresponding to relative humidities between 5 and 45% and F values between 28 and 17. A bivariate fit between the ECHAMP measurements and the concentrations calculated by eq. 1 results in the relation $\text{ECHAMP} = (0.88 \pm 0.02) \times (\text{IU cal source}) + (6.6 \pm 4.5) \text{ ppt}$ with an R^2 of 0.99. If both instrument's calibrations were perfectly accurate, however, the slope would not be expected to equal unity because the two instrument's calibration methods do not produce the same type of peroxy radicals. ECHAMP is calibrated with the acetone photolysis method, which produces an equimolar mixture of CH₃O₂ and CH₃C(O)O₂ radicals (Wood and Charest, 2014). Because a calculated 10% of both of these radicals will be converted to CH₃ONO in the reaction chambers and not be detected, ECHAMP is expected to be 11% (1/0.9) more sensitive to HO₂ than to CH₃O₂ and CH₃C(O)O₂. Moreover, ECHAMP is expected to be between 7 to 12% less sensitive to RO₂ from butane and isoprene than to HO₂ because of the respective alkyl nitrate yields for both peroxy radicals: 8% for butane and 7 – 12% for isoprene (Atkinson et al., 1982; Lockwood et al., 2010; Patchen et al., 2007; Paulot et al., 2009). Thus if both instruments' calibrations were perfectly accurate, then the expected slope for the calibration comparison using butane (i.e., 50% HO₂ and 50% C₄H₉O₂) would be 1.07 (i.e., 1.11×0.96) and the expected slope when using isoprene would be between 1.07 and 1.04 depending on the isoprene alkyl nitrate yield. These values differ from the observed slope of 0.88 by 18 to 22%.

The 2 σ analytical uncertainty for the IU calibration source and ECHAMP measurements are 36% and 27%, respectively. Because the IU calibration source's O₃ mixing ratios were determined by ECHAMP, however, a portion of these two uncertainties is correlated. The uncertainty bars in Fig. 2 have been reduced to remove this component of the uncertainty - to 23% for IU (Dusanter et al., 2008) and 26.6% for ECHAMP. The 18 to 22% difference between the observed slope of 0.88 and the expected slope of 1.04 to 1.07 is within the adjusted uncertainties of both the ECHAMP measurements and the IU calibration source. Moreover, that ECHAMP evidently has near identical sensitivity to these two types of organic peroxy radicals demonstrates that differences in the mechanisms for converting RO₂ to HO₂ between β -hydroxy and alkyl peroxy radicals do not appear to affect their detection by ECHAMP.

The excellent linearity of Fig. 2 is notable because the calibrations were performed over a range of relative humidity values, each of which requires a different amplification factor to be used by ECHAMP. If the RH-dependence of the ECHAMP calibration had been ignored and only the dry calibration factor been used instead, the comparison would have been inferior as indicated by the squares in Fig. 2, for which a linear fit (not shown) gives the relation $\text{ECHAMP} = 0.69(\text{IU cal source}) + 10.8 \text{ ppt}$. This serves as evidence that RH-dependent calibrations are indeed needed for producing accurate results from chemical amplifiers, including traditional CO and NO-based instruments (e.g., PERCA).

3.2 Ambient concentrations of total peroxy (XO₂) radicals, trace gases, and meteorological parameters

Ambient concentrations (15-minute averages) of XO₂, isoprene, ethene, O₃, NO, and NO₂, along with meteorological parameters are shown in Fig. 3.



The 15-min average XO_2 concentrations in the daytime ranged from below the detection limit of ~ 5 ppt to 77 ppt. Among the VOCs measured, the daytime concentrations of low-molecular weight total alkanes ($\text{C}_2\text{-C}_5$) were the highest (average mixing ratio ± 1 standard deviation: 5.7 ± 3.9 ppb) followed by isoprene (3.6 ± 1.9 ppb), total $\text{C}_2\text{-C}_5$ alkenes (1.1 ± 0.3 ppb), high-molecular-weight alkanes ($\text{C}_6\text{-C}_{14}$, 0.3 ± 0.2 ppb), toluene (0.1 ± 0.1 ppb) and monoterpenes (0.1 ppb). NO concentrations typically peaked at 0.2 to 0.8 ppb between 09:00 - 11:00 and were almost always below 0.2 ppb between 12:00 and 21:00, whereas NO_2 concentrations in the daytime ranged between 0.3 to 3 ppb. O_3 concentrations varied between 0 to 71 ppb (av. 35.0 ± 8.4 ppb).

Measured XO_2 concentrations during IRRONIC exhibited a diurnal profile characterized by low mixing ratios (often below detection limit) between 0:00 – 07:00, increasing values from 07:00 to 13:00, peak values between 13:00 and 16:00, followed by a decrease in the late afternoon, similar to past measurements in other forests ([Burkert et al., 2001](#); [Hewitt et al., 2010](#); [Mihele and Hastie, 2003](#)). XO_2 mixing ratios were generally positively correlated with concentrations of isoprene, total alkenes, and ozone (Fig. 3). The highest XO_2 concentrations of over 60 ppt were measured during the afternoon of 24 and 25 July, coinciding with the highest average concentrations of isoprene (4.4 ppb), total alkenes (1.8 ppb), and O_3 (61 ppb), and the lowest average concentration of NO (0.1 ppb). The lowest daytime concentrations of XO_2 were observed on 13 July and 15 July, which were also characterized by lower isoprene and ozone mixing ratios and higher NO_2 mixing ratios.

We compare our XO_2 concentrations with reported XO_2 and HO_2^* concentrations from other forests. The observed daytime XO_2 mixing ratios (campaign daytime average 26 ppt) at the IRRONIC site at Indiana are similar to those reported in a tropical rain forest in Malaysia (range 2-68 ppt) ([Hewitt et al., 2010](#)), in a northern Michigan forest during several intensive campaigns (range 8-65 ppt) ([Griffith et al., 2013](#); [Mihele and Hastie, 2003](#)), and in a tropical forest over South America (campaign av. 42 ppt) ([Lelieveld et al., 2008](#)). XO_2 concentrations at Indiana never exceeded 80 ppt, in contrast to studies in which measured peroxy radical mixing ratios sometimes exceeded 150 ppt ([Burkert et al., 2001](#); [Wolfe et al., 2014](#)).

Measurements of peroxy radical and NO concentrations enable ozone production rates to be calculated directly rather than rely on photochemical models. Using the measured concentrations of peroxy radicals and NO, calculated ozone production rates at the IURTP were at most 9 ppb/hr and described more in the SI.

3.3 Comparisons of Ambient Peroxy Radical Mixing Ratios

Figure 4 compares ambient $[\text{XO}_2]$ measurements by ECHAMP (30-minute averages) with the $[\text{HO}_2^*]$ measurements by LIF-FAGE (1-minute average every 30 minutes) during 13-25 July. Only data from days in which both instruments were operational are shown. Although in general it is preferable to compare measurements with equal time averaging, the precision of ECHAMP during this campaign – typically 2.5 ppt (1σ) for the 1.5 minute average measurements – necessitated this averaging. The diurnal profiles of both measurement sets, divided into 30-minute bins, are displayed in



Fig. 5. Both figures indicate that the ECHAMP and LIF-FAGE measurements are in general well correlated and follow the same diurnal trend, though closer inspection reveals significant day to day and even hour to hour variability in the ratio.

A linear regression of all the XO_2 and HO_2^* measurements yields the relationship $[\text{XO}_2] = 0.85 [\text{HO}_2^*] + 3.3 \text{ ppt}$, $R^2 = 0.67$ (Fig 6.). Adjusting this for the calibration difference described in section 3.1, i.e. increasing the XO_2 values by a factor of 1.2 yields a slope of 1.02. If the y-intercept is constrained to zero, which is not necessarily appropriate, the slope increases to 0.96, or 1.15 after adjusting for the calibration difference. The individual days with the highest $[\text{XO}_2]/[\text{HO}_2^*]$ slopes were 24 and 25 July, with slopes of 1.25 and 1.08, respectively (1.5 and 1.3 after adjusting for the calibration difference). These two days were characterized by the highest mixing ratios of O_3 , isoprene, and the anthropogenic VOCs ethene and ethyne. This linear fit of all the data is difficult to interpret, however, since the XO_2 measurements are 30 minute averages and the HO_2^* measurements are 1-minute averages taken every 30 minutes. A regression of the binned data shown in Fig. 5 gives the relation $[\text{XO}_2] = 0.98 [\text{HO}_2^*] + 1.7 \text{ ppt}$; accounting for the calibration difference gives an adjusted slope of 1.18.

To further investigate the effect of this different averaging on the comparison, on 22 July the IU-LIF-FAGE instrument operated in HO_2^* -only mode (i.e., with no time devoted to measuring OH). We compare the resulting 1-minute and 15-minute averaged HO_2^* measurements to the 1.5 minute and 15-minute averaged XO_2 measurements (Fig. 7). Between 15:00 and 17:00, the HO_2^* measurements increased from 50 to 70 pptv and decreased back to 50 pptv while the XO_2 measurements were relatively invariant at 40 pptv. Ignoring the difference between the average mixing ratios, this difference in the temporal profile of the two instruments' measurements result could only be "real" if there were changes in the peroxy radical relative composition on this two-hour time scale, e.g. a simultaneous increase in HO_2 and a decrease in alkyl peroxy radicals, such that $[\text{HO}_2^*]$ actually did increase while the mixing ratio of total peroxy radicals was almost constant. Measurements of VOC composition and NO_x do not support such a fast change in peroxy radical composition, suggesting that these observations were more likely the result of an instrumental issue. Currently we are unable to identify the exact cause of this observation, but possible explanations are a transient interference in the HO_2^* measurement when sampling ambient air or a change in the sensitivity of the ECHAMP measurements.

The "true" $[\text{XO}_2]/[\text{HO}_2^*]$ ratio, i.e., the ratio that would be produced by the two instruments' measurements if they were calibrated to the same source and operated exactly as expected without any uncharacterized interferences or losses, depends on the composition of the peroxy radicals. As described in Section 2 (Experimental Methods), for both ECHAMP and LIF-FAGE, the sensitivity of the instrument to individual RO_2 compounds depends on the R-group and is characterized by the parameter " α ", which is the instrument's sensitivity to each RO_2 relative to its sensitivity to HO_2 . For ECHAMP α is determined largely by the fraction of RO_2 that is converted to alkyl nitrates (RONO_2) and alkyl nitrites (RONO) following reaction with NO at atmospheric pressure. For LIF-FAGE, α is mostly determined by how quickly each RO_2 is converted sequentially to HO_2 and then OH following reaction with NO after the expansion of the sampled gas into the low-pressure region of the instrument (Fuchs et al., 2011; Lew et al., 2018). Air in which CH_3O_2 , $\text{CH}_3\text{C}(\text{O})\text{O}_2$, and small (<C5) alkyl peroxy radicals have a large contribution to the total peroxy radical concentration would thus produce a relatively high



$[XO_2]/[HO_2^*]$ value, since ECHAMP is sensitive to those peroxy radicals ($\alpha > 0.9$) whereas the LIF-FAGE HO_2^* measurement is not ($\alpha < 0.1$). In contrast, air with a relatively high fraction of alkene-derived RO_2 (e.g., isoprene peroxy radicals), for which both ECHAMP and LIF-FAGE HO_2^* α values are near one, would be expected to lead to lower $[XO_2]/[HO_2^*]$ values (i.e., closer to unity).

5 Because the composition of the peroxy radicals during IRRONIC is not exactly known, we examine the predicted speciation generated by zero-dimensional photochemical modeling of the IRRONIC dataset using two versions of the Regional Atmospheric Chemistry Mechanism (RACM2 and RACM2-LIM1) and the Master Chemical Mechanism (MCM 3.2 and 3.3.1). A full comparison of the modeled and measured concentrations is beyond the scope of this paper; we use these model outputs mainly to inform the discussion of the relative speciation of total peroxy radicals and its relation to the
10 expected and measured $[XO_2]/[HO_2^*]$ ratio. A fuller description of the photochemistry at this site, including OH reactivity measurements, will be described in a companion paper (Lew et al, in preparation).

The accuracy of the model results is, of course, subject to how comprehensive and accurate the supporting measurements and underlying chemical mechanisms are, but nonetheless help to frame the interpretation of the two instruments' measurements. Due to gaps in the NO data because of problems with the Thermo chemiluminescence sensor,
15 there are only three days for which we have model results and measured peroxy radical concentrations by both ECHAMP and LIF-FAGE – on the 16th, 22nd, and 24th of July. The model was run for these three days, and also a diurnal profile for the entire campaign was run using diurnal average concentrations of constrained species. From these model results we calculate the expected values measured by ECHAMP and LIF-FAGE based on each instrument's relevant values for α :

$$20 \text{ ECHAMP } [XO_2]_{\text{EXPECTED}} = [HO_2] + 0.9([CH_3O_2]) + 0.92([C_5H_8(OH)O_2]) + 0.9([CH_3C(O)O_2]) + 0.9(\text{Other}) \quad (4)$$

$$\text{LIF-FAGE } [HO_2^*]_{\text{EXPECTED}} = [HO_2] + 0.05([CH_3O_2]) + 0.83([C_5H_8(OH)O_2]) + 0.05([CH_3C(O)O_2]) + 0.7(\text{Other}) \quad (5)$$

The “Other” category includes all types of peroxy radicals, e.g., from monoterpenes, ethene, etc.

25 The top portion of Fig. 5 shows the average diurnal profile for the modeled and measured $[XO_2]/[HO_2^*]$ ratio using all days when there were both XO_2 and HO_2^* measurements. Between 10:00 and 18:00 the modeled $[XO_2]/[HO_2^*]$ ratio using RACM2 varied between 1.2 and 1.4, whereas the measured ratio varied between 0.9 and 1.4, with a greater amount of variability from hour to hour. Increasing the observed ratio by 20% to account for the calibration comparison (section 3.1) gives an adjusted measured ratio of between 1.1 and 1.7. The highly variable ratios during nighttime mainly reflect the lower
30 signal to noise ratios of both instruments when peroxy radical concentrations were low (less than ~5 ppt).

Measured and RACM2 modeled concentrations for 16, 22, and 24 July are shown in Fig. 8. On all three days the relative contributions from the various types of peroxy radicals are comparable. At 15:30 –when concentrations are highest – the modeled peroxy radicals comprised 40% $C_5H_8(OH)O_2$, 33% HO_2 , 16% CH_3O_2 , 4% $CH_3C(O)O_2$, and 7% “Other”, resulting in an expected (modeled) $[XO_2]/[HO_2^*]$ ratio of 1.3. The measured $[XO_2]/[HO_2^*]$ ratio is close to unity on 16 and



22 July, and between 1.2 and 1.5 on 24 July. Increasing these measured ratios by 20% to account for the calibration comparison results produces adjusted measured $[XO_2]/[HO_2^*]$ ratios of 1.2 on 16 and 22 July and 1.4 to 1.8 on 24 July.

Measured $[XO_2]$ mixing ratios are ~40% lower than the modeled $[XO_2]$ on 16 and 22 July but within 10% on 24 July. The comparison between measured $[HO_2^*]$ and modeled $[HO_2^*]$ for these three days exhibits more variability (Fig. 8), but except for a few short time periods modeled $[HO_2^*]$ is greater than measured $[HO_2^*]$. Choice of chemical mechanism does lead to differences in both the absolute concentrations and relative speciation. Compared to RACM2, RACM2-LIM1 leads to 20% higher peak concentrations though the relative speciation is very similar. Concentrations produced by MCM v3.2 are similarly 20% higher than RACM2, though $CH_3C(O)O_2$ accounts for 9% of the total peroxy radical concentration in the afternoon compared to 4% for RACM2. MCM v3.3.1 leads to the highest overall concentrations (~30% higher than RACM2), a 12% contribution from $CH_3C(O)O_2$, and only a 19% contribution from $C_5H_8(OH)O_2$. Further details can be found in the SI. In summary, the $[XO_2]/[HO_2^*]$ ratios modeled by the four chemical mechanisms vary from 1.3 (RACM2 and RACM2-LIM1) to 1.4 (MCM v. 32) to 1.5 (MCM v. 331) between 15:00 and 16:00, whereas the measured ratios varied from 0.8 to 1.3 depending on the day, with an average value of 1.0. After accounting for the 20% calibration difference, the modeled and measured ratios agree to within the experimental and model uncertainties. That both the measured $[XO_2]$ and $[HO_2^*]$ concentrations, however, are typically only 50% of the values predicted by MCM v3.3.1 does suggest that the model somehow over predicts peroxy radical concentrations, in agreement with past results from high isoprene environments (Griffith et al., 2013; Whalley et al., 2011; Mao et al., 2012).

Observations of $[XO_2]/[HO_2^*]$ ratios less than one were observed during parts of 13, 17, and 18 July and even after increasing by 20% to account for the calibration comparison do not seem reasonable or in some cases even possible. These observations were most likely caused by issues with one or both instruments. Two possible causes that warrant investigation in subsequent field measurements are discussed below:

1. *Error in the ECHAMP calibration, especially for RH values greater than 45%.* Although the calibration comparison presented in section 3.1 show that the ECHAMP and LIF-FAGE instrument's calibrations agreed to within measurement uncertainties, that is not necessarily true for RH values greater than those used during those calibration tests. The highest RH value during the calibration comparisons was 45%, whereas the daytime minimum RH values between 12:00 and 16:00, when measured $[XO_2]$ and $[HO_2^*]$ were both highest, were typically between 45% and 65% (Fig 1). Furthermore, we cannot prove that the ECHAMP calibration was invariant from day to day. We include potential sampling losses to be a part of the overall ECHAMP calibration.

2. *Interferences in the LIF-FAGE measurement.* The comparison of high temporal resolution in Fig. 7 revealed differences in the temporal profile of the LIF-FAGE and ECHAMP sensor. If these were caused by an interference in the LIF-FAGE measurement when sampling ambient air, then it would follow that the two instruments would agree when sampling a calibration source but differ when sampling ambient air.



As discussed earlier, the RH-dependence of the sensitivity of chemical amplifiers has recently been questioned (Andrés-Hernández et al., 2010; Sommariva et al., 2011). Had we ignored the RH dependence for ECHAMP's amplification factor and simply used the value under dry conditions, the daytime XO_2 values would have been roughly 50% lower than those presented in this paper, leading to unrealistically low $[\text{XO}_2]/[\text{HO}_2^*]$ ratios of ~ 0.5 .

4. Conclusions

The results of this comparison of the IU calibration source and the ambient measurements of peroxy radicals by ECHAMP and LIF-FAGE provide encouraging first results that the newly developed ECHAMP method can be used for ambient measurements of total peroxy radicals. The ECHAMP measurements, based on the acetone photolysis method, and the IU water vapor photolysis calibration source agreed within 12%, within the experimental uncertainties. The measured mixing ratios of XO_2 and HO_2^* were lower than the concentrations predicted by the RACM2, RACM2-LIM1, MCM v. 3.2, and MCM v. 3.3.1 chemical mechanisms. The measured $[\text{XO}_2]/[\text{HO}_2^*]$ ratios, however, differed from the ratios predicted by zero-dimensional photochemical modeling by less than the combined measurement and modeling uncertainties.

An attribute of these comparison exercises is that the two instruments operate on very different measurement principles and the calibration methods differ greatly. Although the calibration comparison was favorable, due to the time required to conduct successful calibrations with the acetone photolysis method and its overall inconvenience (Wood and Charest, 2014) we have discontinued its use. For subsequent field measurements we have used the water vapor photolysis method and another method based on methyl iodide photolysis (Anderson et al., 2018; Clemmitshaw et al., 1997; Liu and Zhang, 2014). All three calibration methods do indicate that a humidity-dependent calibration must be used for both CO -based and ethane-based chemical amplifiers.

Data Availability

Data are available upon request from the corresponding author (Ezra.Wood@drexel.edu)

Author contributions.

EW and PS designed the research project. SK and EW were responsible for the ECHAMP measurements and supporting measurements of NO , NO_2 , and O_3 . ML, BB, PR, and PS were responsible for the LIF-FAGE measurements and photochemical modeling. SD, SS, TL, and NL were responsible for the measurements of VOCs. SK and EW conducted the analysis and wrote the paper with feedback from all co-authors.



Competing Interests. The authors declare that they have no conflicts of interest.

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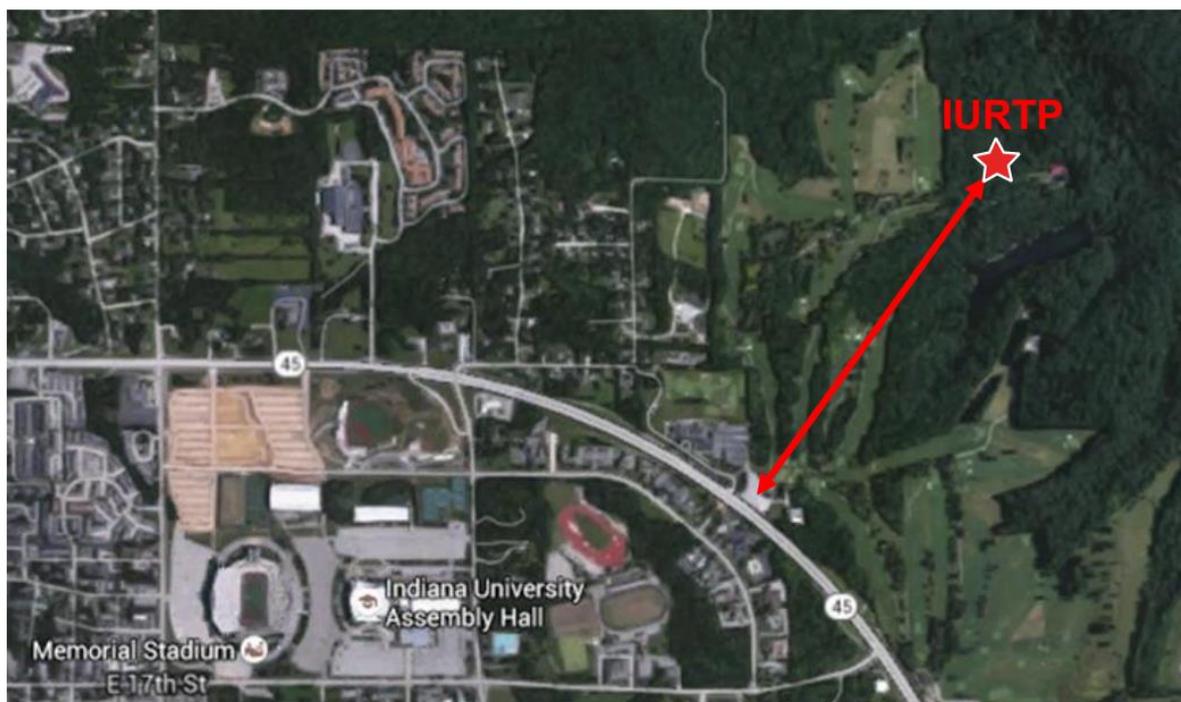
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5 **Figure 1.** Map of the sampling site. The star symbol represents the Indiana University Research and Teaching Preserve (IURTP) in Bloomington, Indiana, USA.

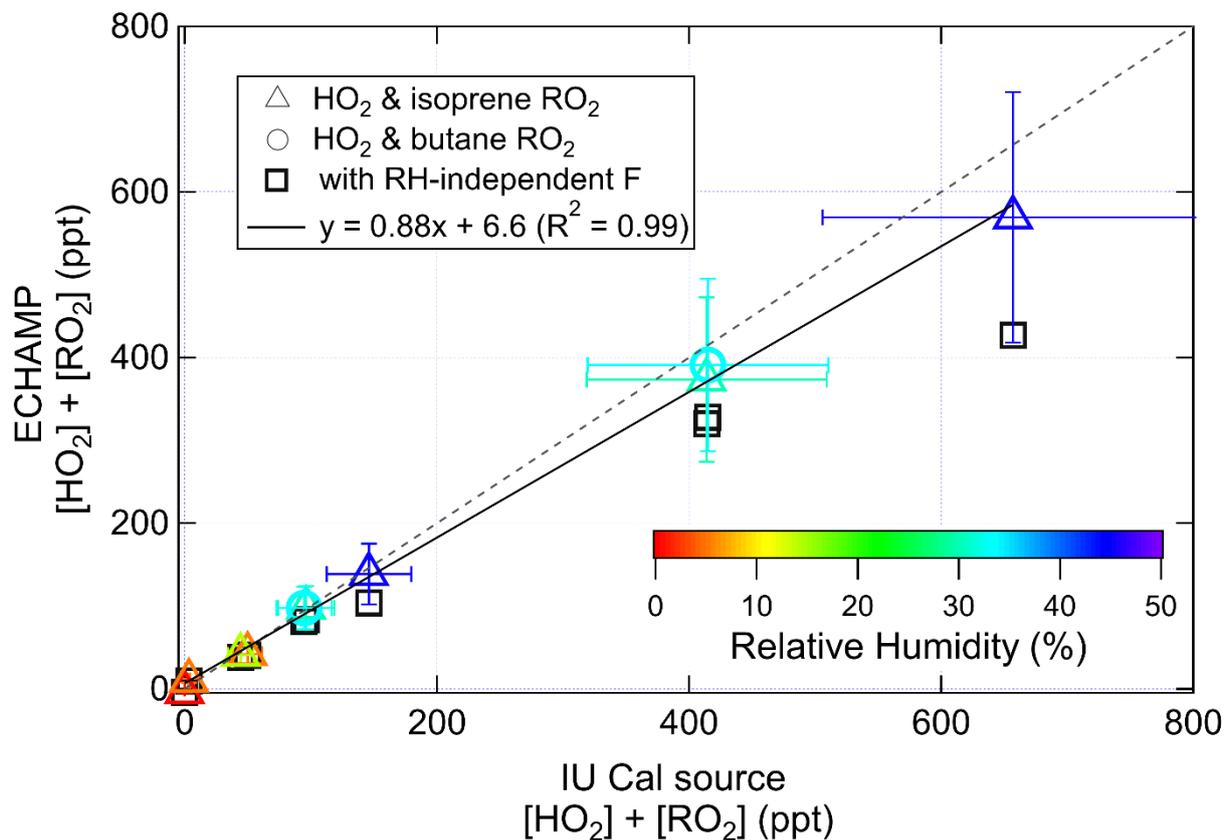


Figure 2. Results of the calibration comparison in which ECHAMP measured the total peroxy radical concentration in the output of the IU calibration source. The error bars indicate 2σ uncertainties of the ECHAMP measurements and IU calibration source, adjusted for the fact that the IU actinometry was based on the ECHAMP NO_2 calibration. The slope of the dotted line is unity.

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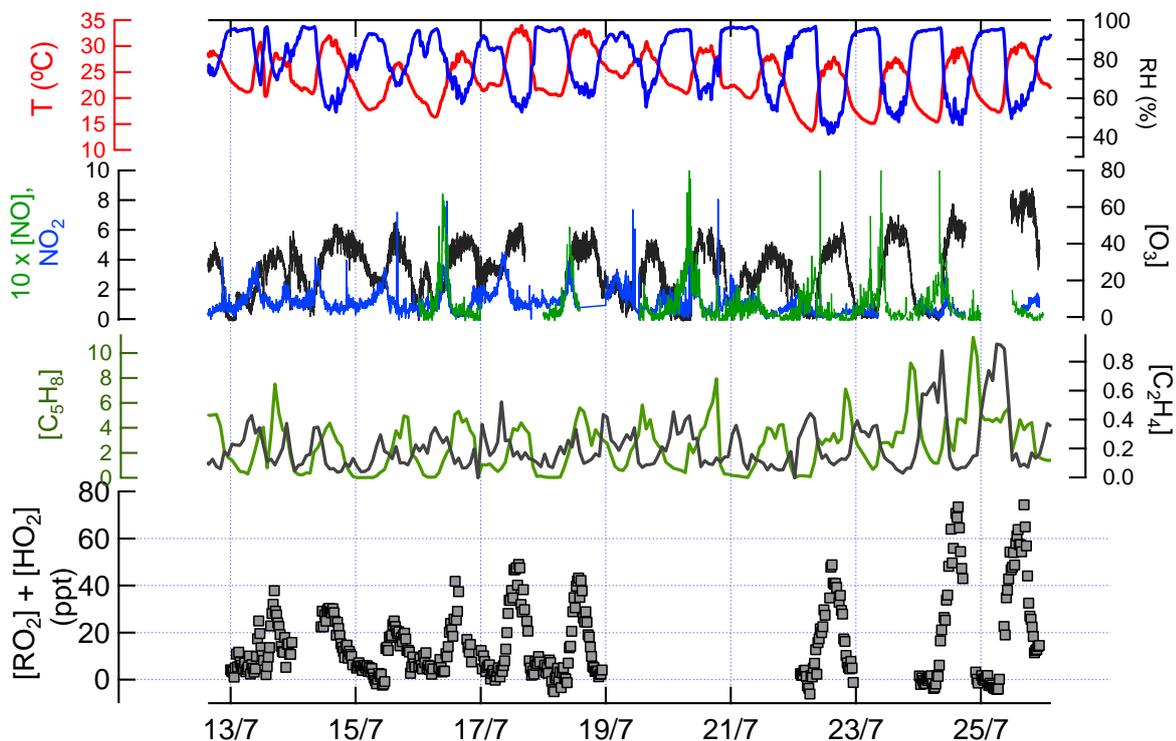


Figure 3. Time series data of measured chemical and physical parameters during IRRONIC. Except where noted, all measurements are in ppb.

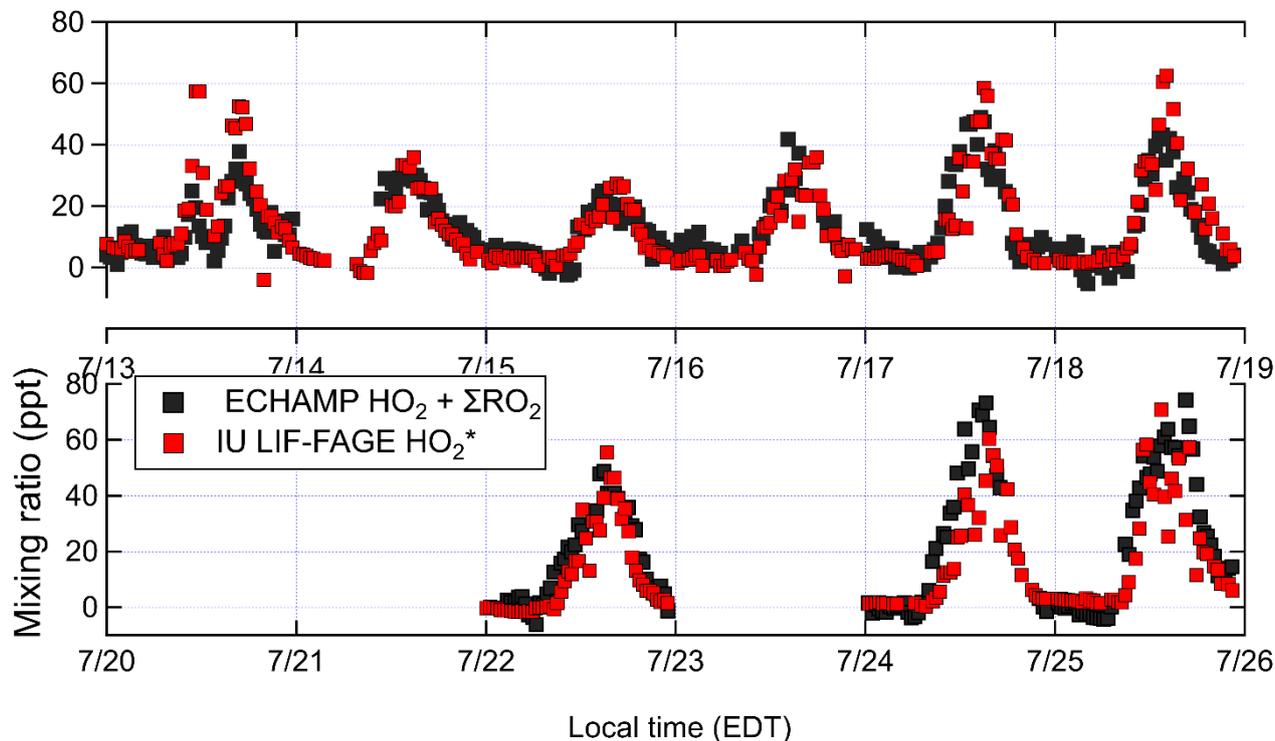
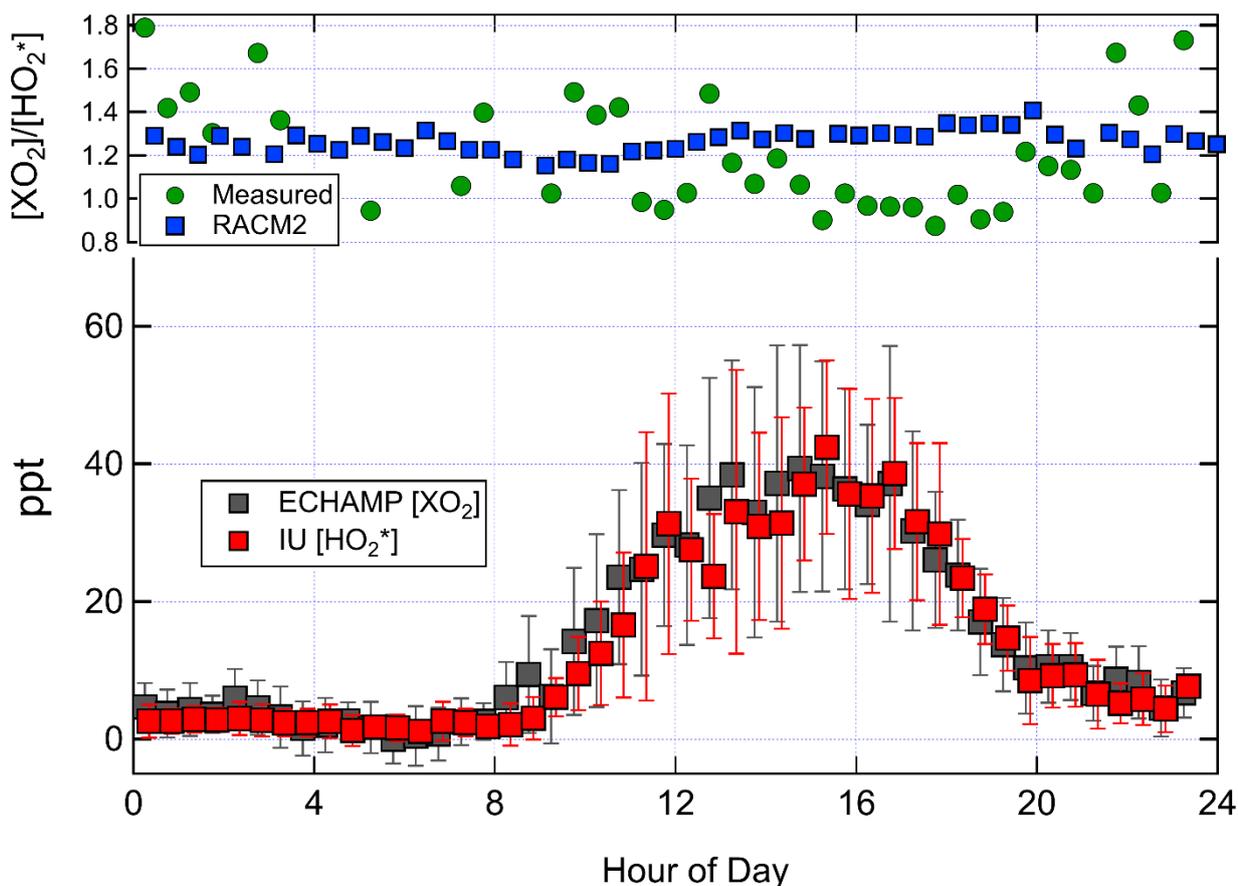


Figure 4. Concentrations of ambient total peroxy radicals (XO_2) by ECHAMP and HO_2^* by IU-LIF-FAGE. 30-minute averaged measurements are shown for ECHAMP XO_2 . For HO_2^* , measurements are 1-minute averages every 30 minutes.



5 **Figure 5. Lower plot: Mean diurnal profile of ECHAMP XO_2 and IU-LIF-FAGE HO_2^* measurements for the 9 days in which both instruments were operational. The HO_2^* values are displayed with a 6 minute horizontal offset for clarity. The error bars indicate the $\pm 1\sigma$ precision of the measurements in each 30-minute time bin. The upper plot shows the $[XO_2]/[HO_2^*]$ ratio - both measured by the two instruments and modeled using the RACM2 chemical mechanism.**

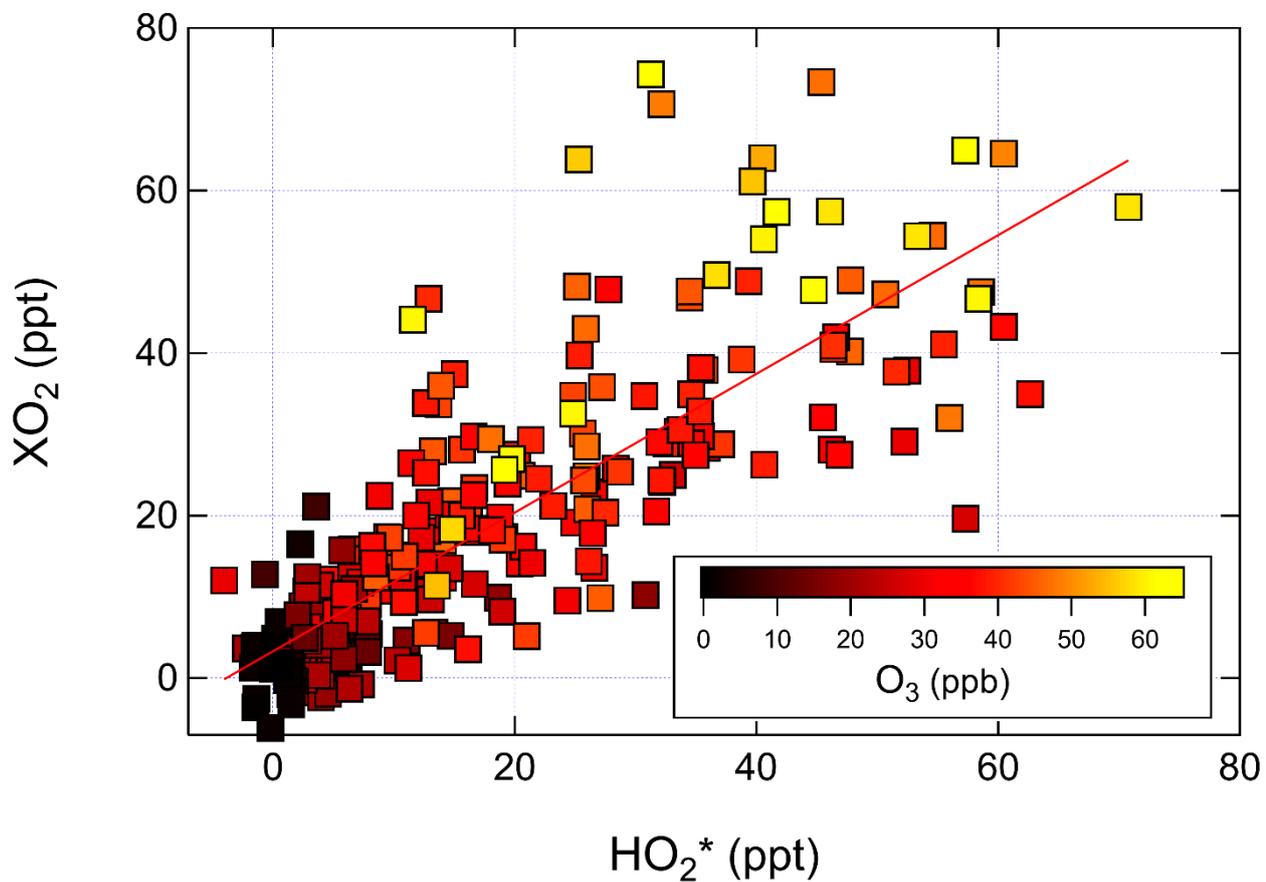


Figure 6. Correlation of ambient $[XO_2]$ measured by ECHAMP with $[HO_2^*]$ measured by IU-LIF-FAGE.

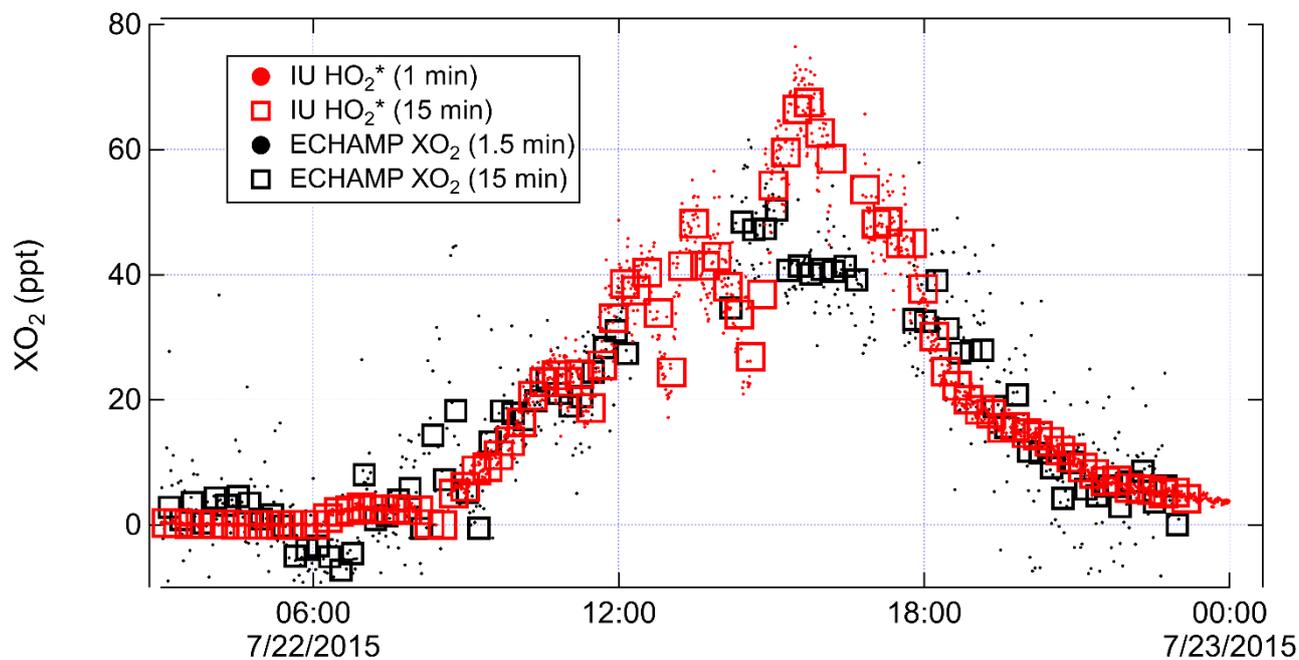


Figure 7. Time series comparing IU LIF-FAGE HO₂* and ECHAMP XO₂ measurements from 22 July, 2015 when the IU LIF-FAGE instrument was run in HO₂*-only mode.

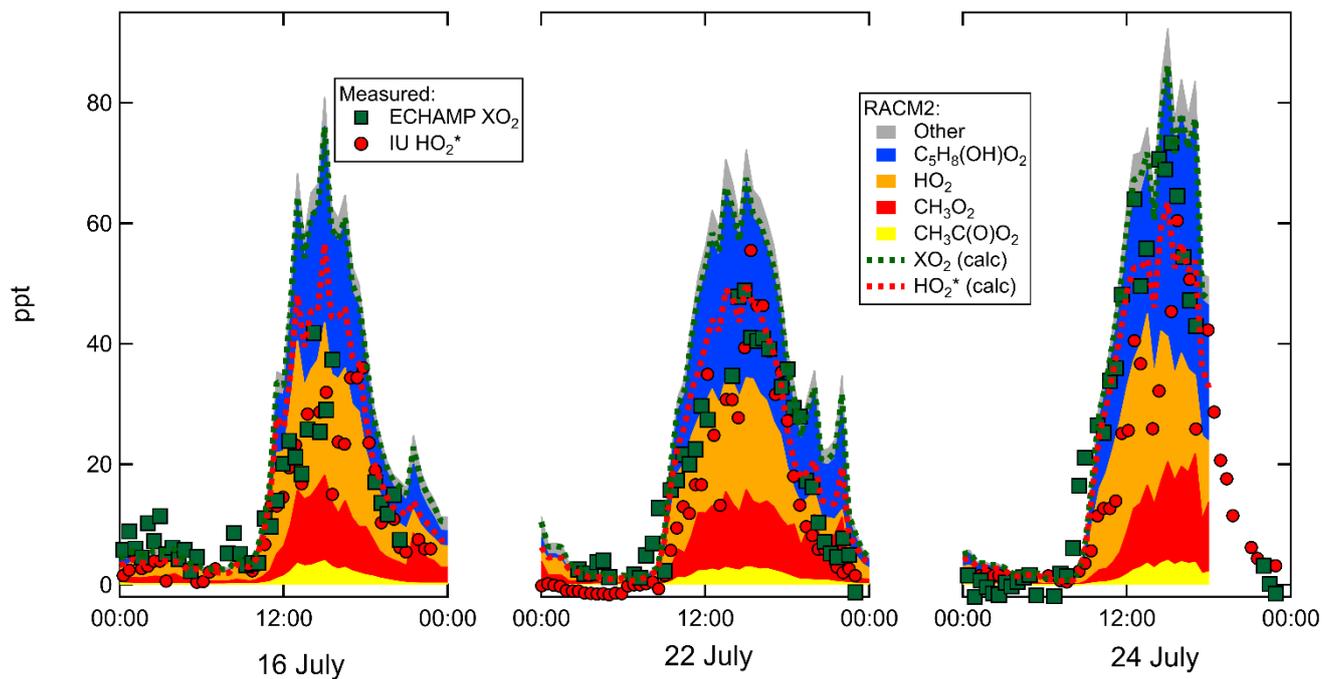


Figure 8. Peroxy radical mixing ratios measured by ECHAMP and LIF-FAGE and modeled by RACM2.

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